- Instruments:
  - HALOE
    - 1991-2005
  - POAM III
    - 1998-2005
  - WVMS
    - 1992-present
  - AURA MLS
    - 2004-present





## Water Vapor Mm-wave Spectrometer (WVMS)



22 GHz radiometers using pressure broadening information to make water vapor profile measurements from 40-80 km WVMS measurements are made at 3 sites of the Network for the Detection of Stratospheric Change (NDSC). Two of these sites are currently operating: Lauder, New Zealand (45°S, 169.7°E): Nov. 1992-Apr. 1993, Jan. 1994-present Mauna Loa, HA (19.5°N, 204.4°E): Mar. 1996-present







## Water Vapor Mm-wave Spectrometer (WVMS)

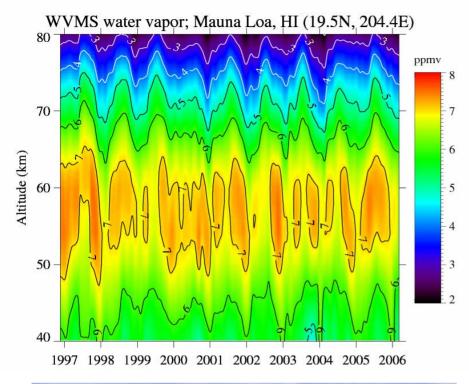


Status of instrument at Table Mountain California (34.4°N, 242.3°E):

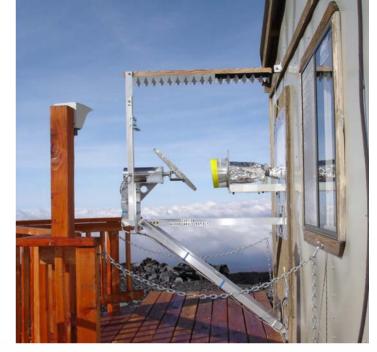
- Data in NDSC database for May 1993-Nov. 1997
- Data available from Nov. 2003-July 2005, but this data should not be compared to the 1993-1997 data.
- Instrument is currently not operating
  - Problems with conversion from DOS to Windows operating system
  - Hardware tests implemented to help determine cause of differences between 1993-1997 and 2003-2005 data.



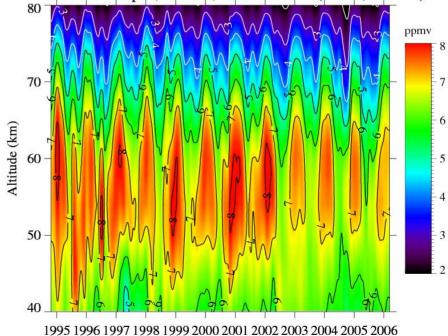








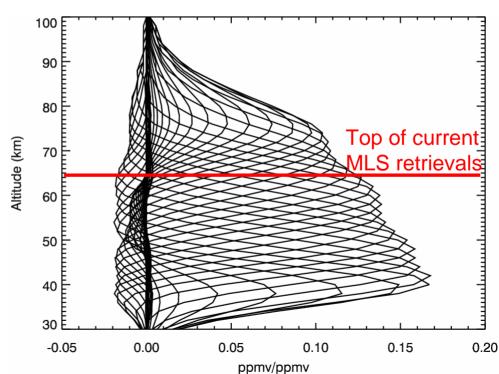
WVMS water vapor; Lauder, New Zealand (45.0S, 169.7E)



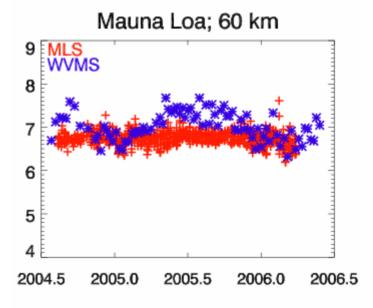
# **Vertical Resolution**

- Vertical resolution of MLS AURA water vapor is estimated to be:
  - 3 km in the upper troposphere and lower stratosphere
  - 4 km in most of the stratosphere
  - 6 km in the lower mesosphere
- Vertical resolution of POAM water vapor is estimated to be:
  - 1 km up to ~30 km
  - Degrading with altitude from 1 to 2 km from 30-40 km
- Vertical resolution of WVMS:

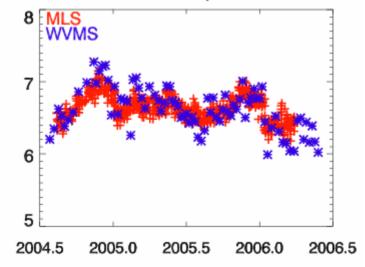
## 1 week integration WVMS3, Mauna Loa



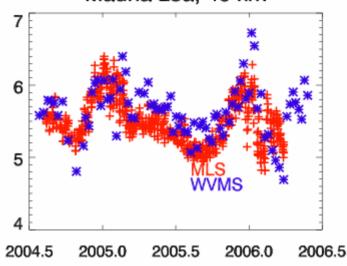
# WVMS-MLS (convolved) comparisons

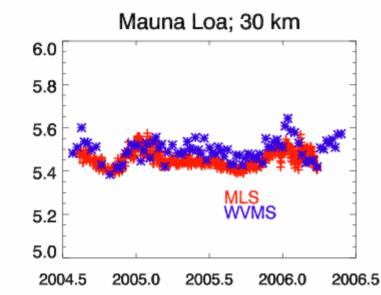


Mauna Loa; 50 km

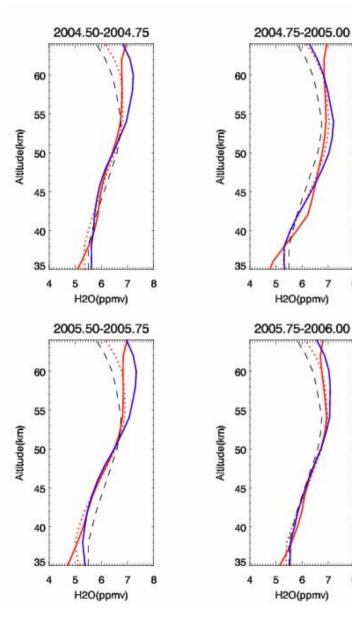


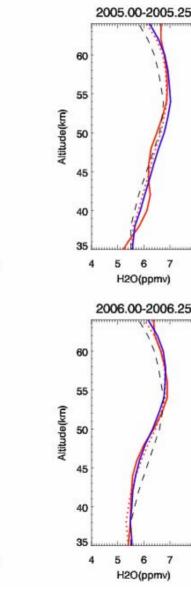
Mauna Loa; 40 km





### WVMS Mauna Loa (19.5N, 204.4E) MLS +/-2 deg latitude and 30 deg longitude





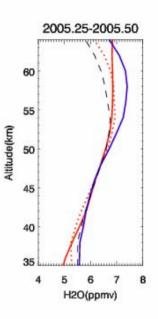
8

8

6 7

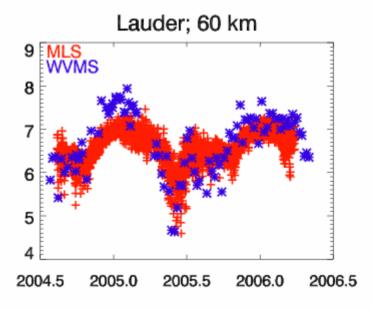
6 7 8

8

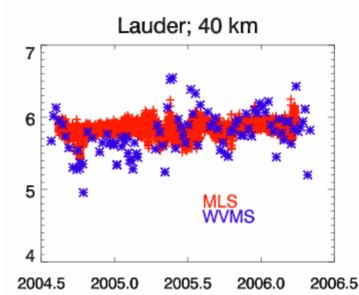


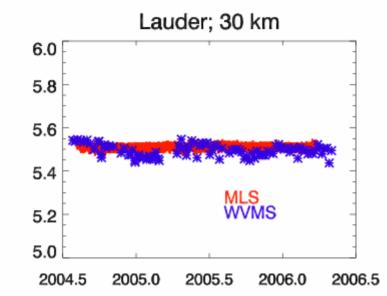
Solid red= **Unconvolved MLS** Solid blue= **WVMS** Dotted red= **Convolved MLS** Dashed= WVMS a priori

# WVMS-MLS (convolved) comparisons

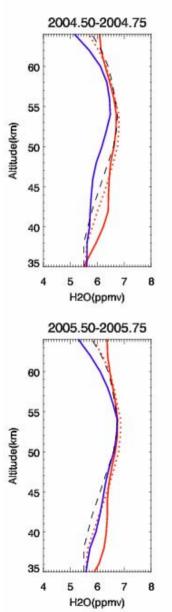


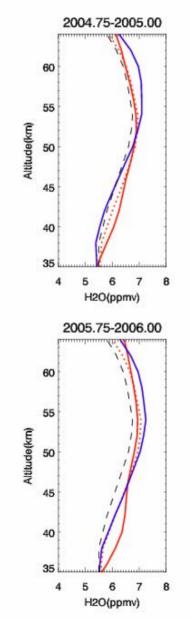
Lauder; 50 km

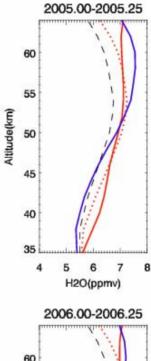


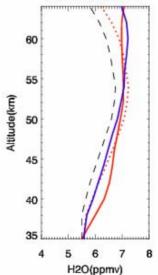


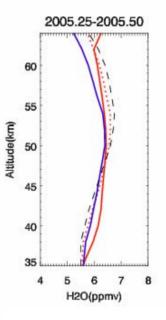
### WVMS Lauder (45.0S, 169.7E) MLS +/-2 deg latitude and 30 deg longitude





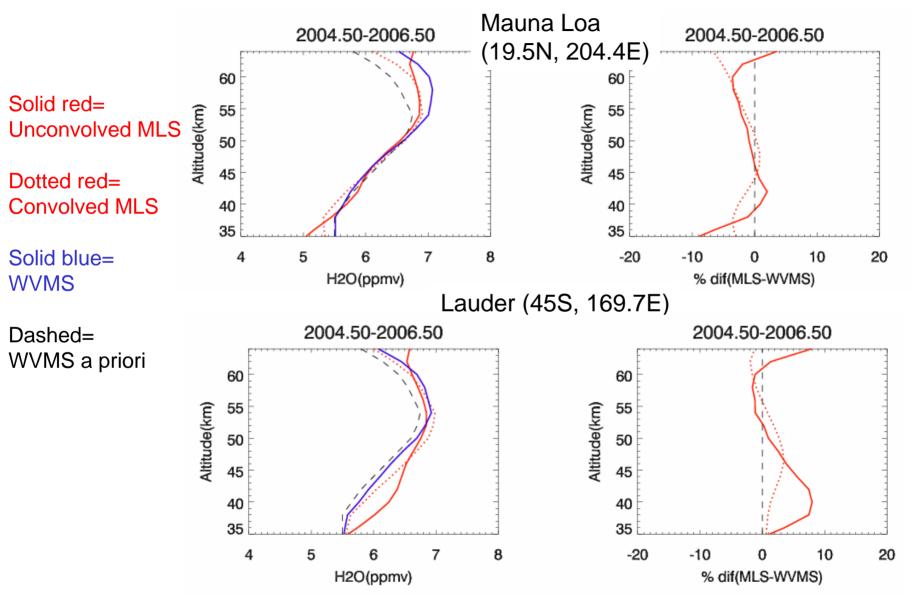




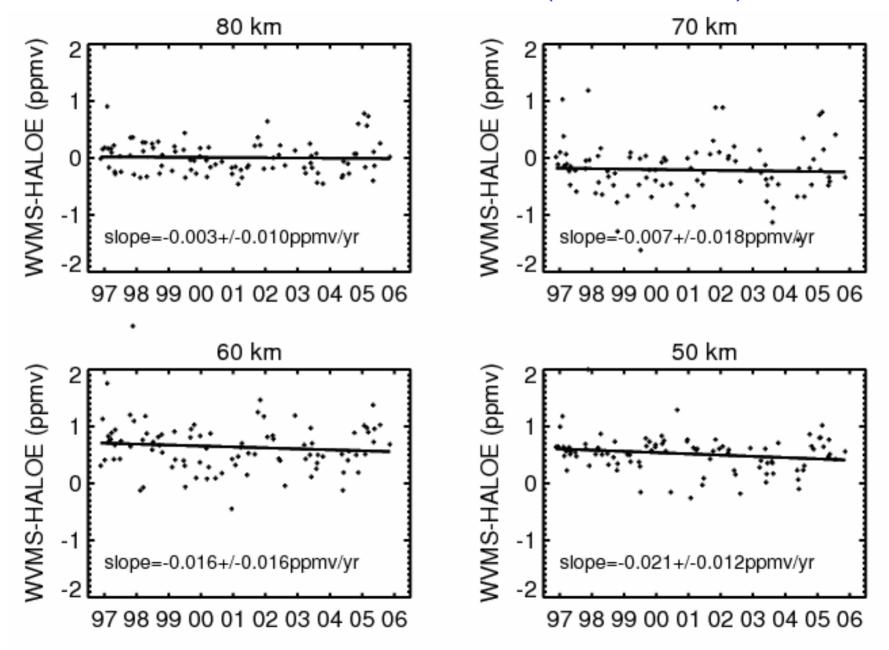


Solid red= Unconvolved MLS Solid blue= WVMS Dotted red= Convolved MLS Dashed= WVMS a priori

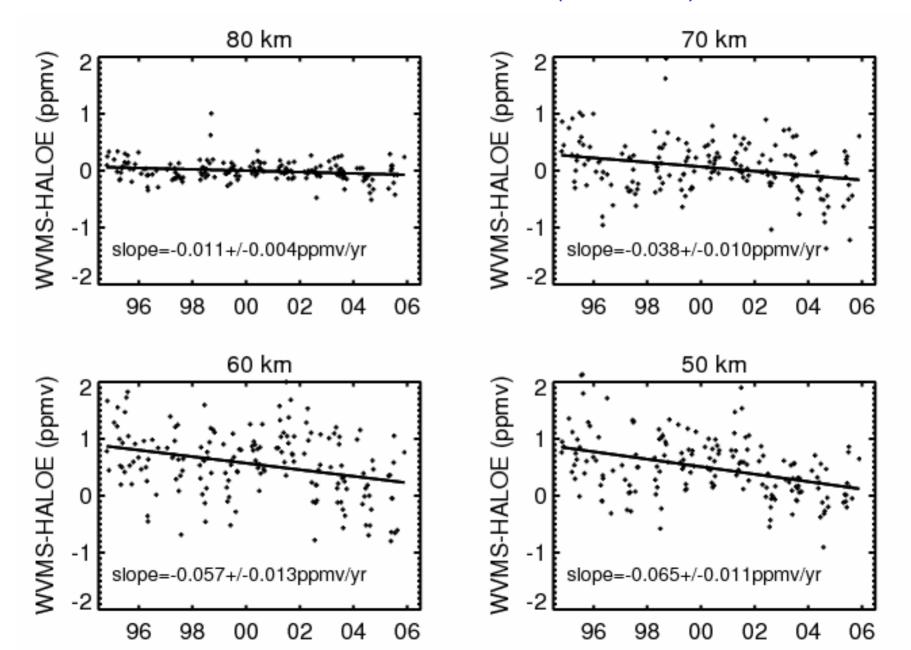
# MLS-WVMS comparisons Averages



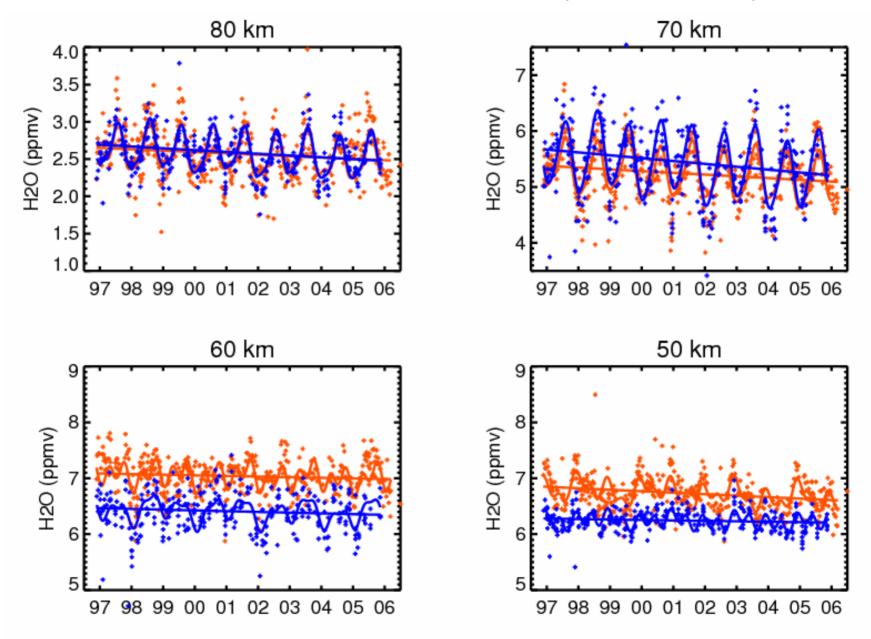
#### WVMS at MLO - HALOE (14.5N-24.5N)



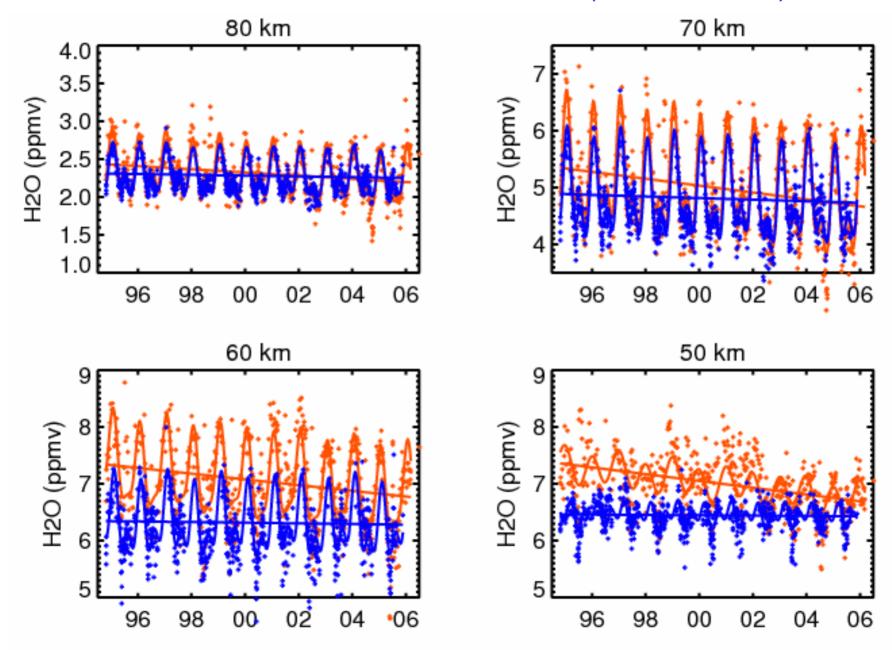
#### WVMS at Lauder - HALOE (40S-50S)



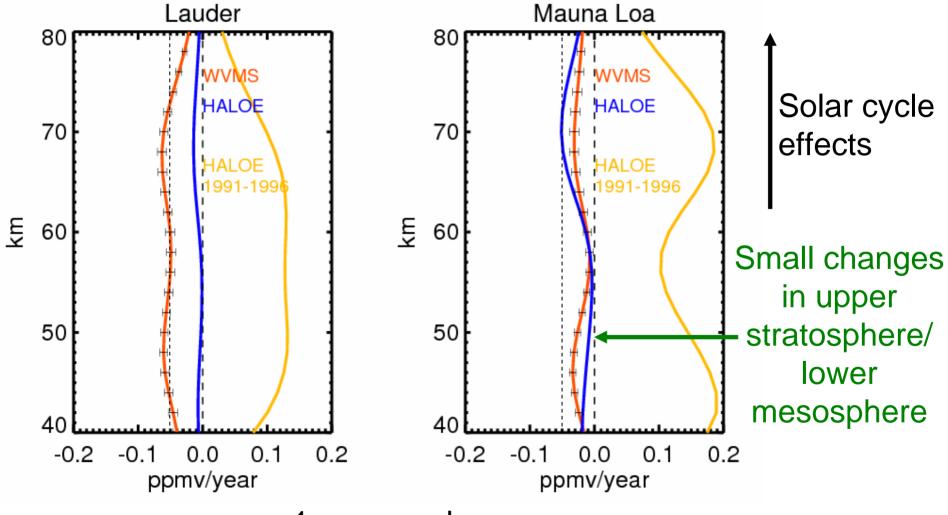
### WVMS at MLO and HALOE (14.5N-24.5N)



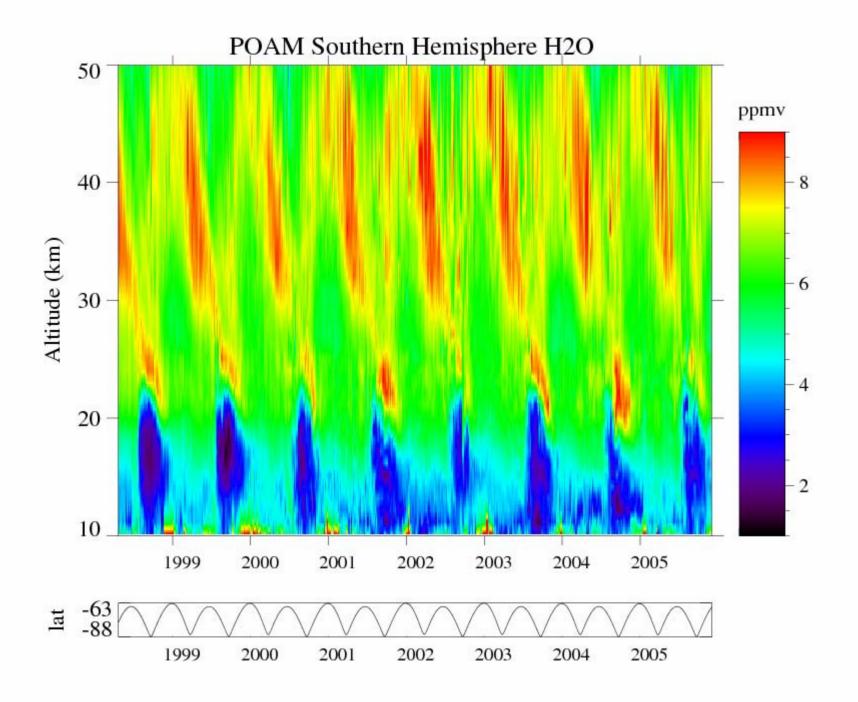
### WVMS at MLO and HALOE (14.5N-24.5N)



# Lauder (45S) and Mauna Loa (14.5N) trends since November 1996

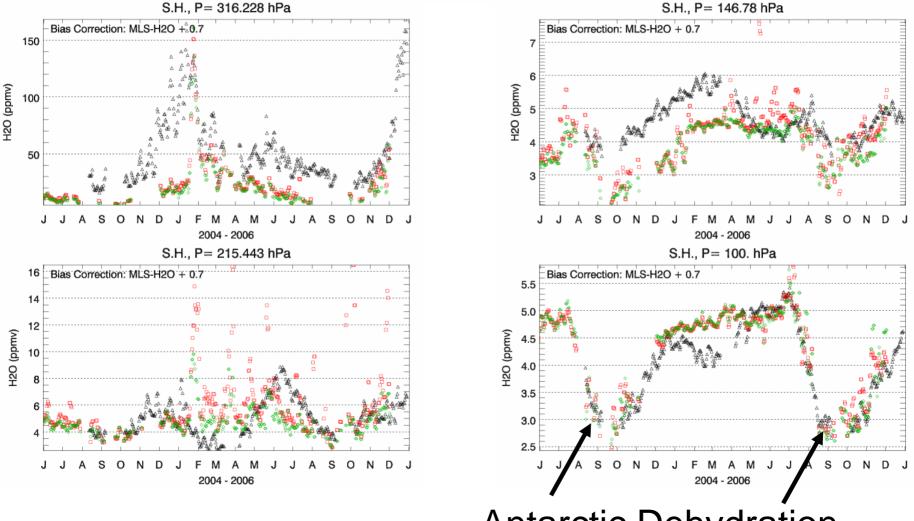


 $1\sigma$  errors shown



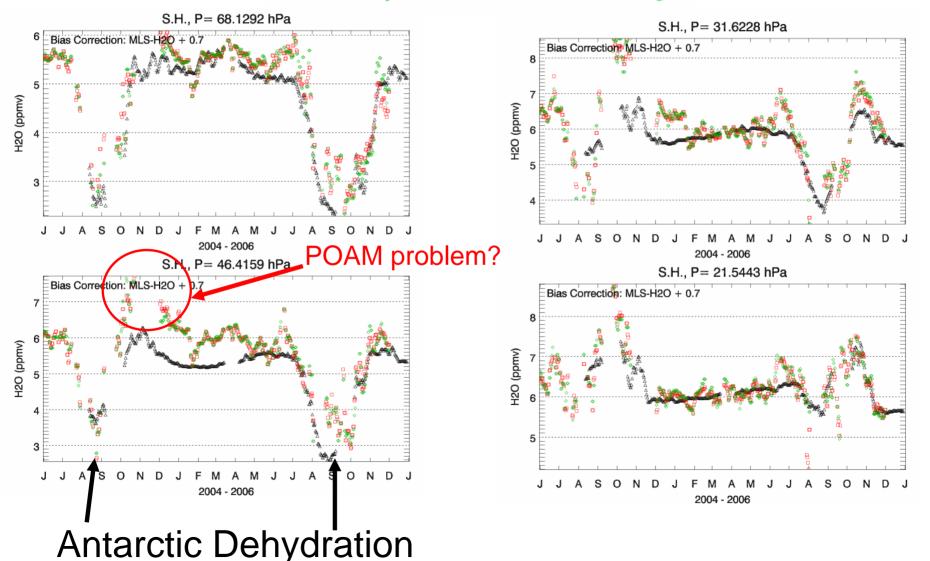
## Troposphere/lowermost stratosphere measurements

MLS 5 Day Ave. at +/- 2° of POAM Lat. POAM 5 Day Average, 3 km vertical smoothing POAM 5 Day Median, 3 km vertical smoothing



Antarctic Dehydration

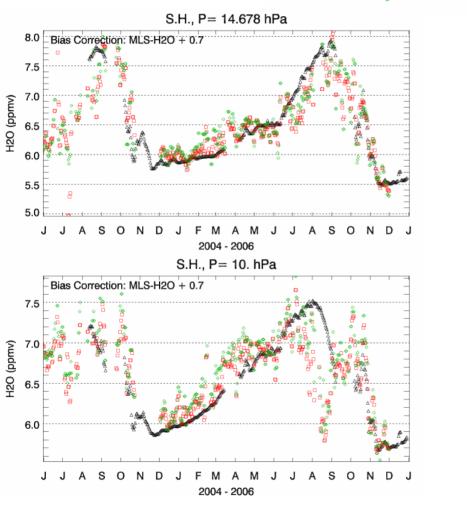
### Lower stratosphere measurements

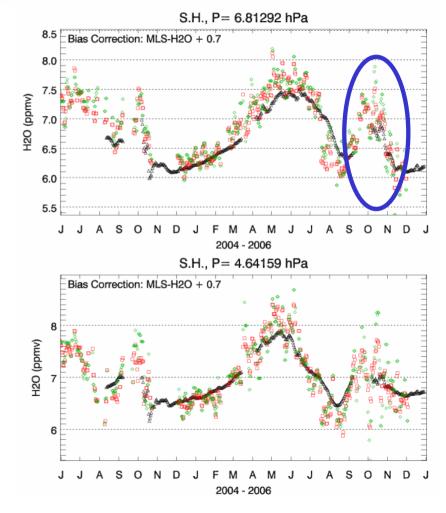


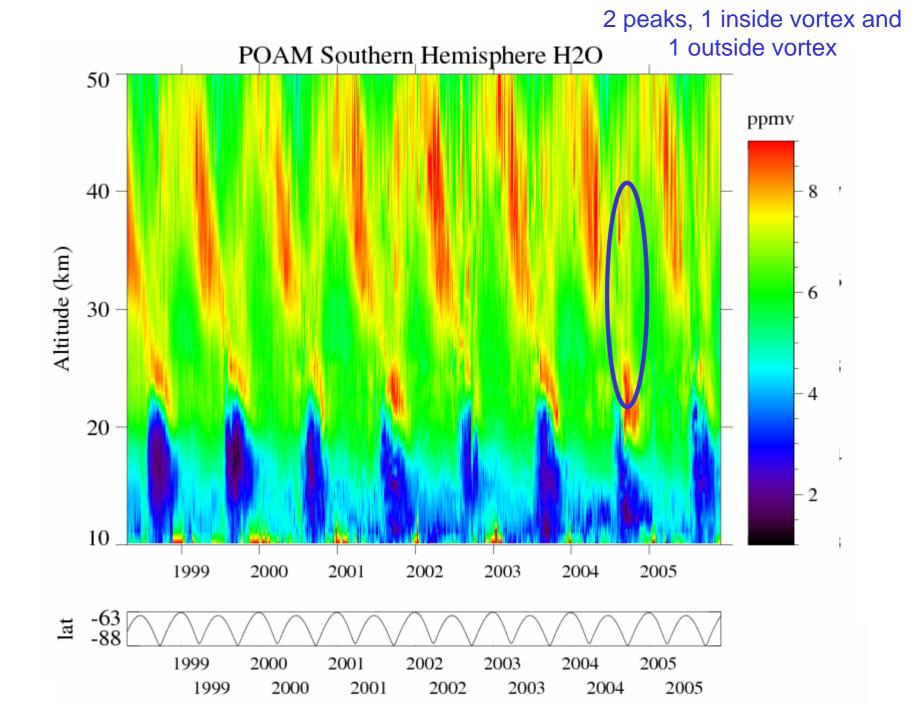
### Upper stratosphere measurements

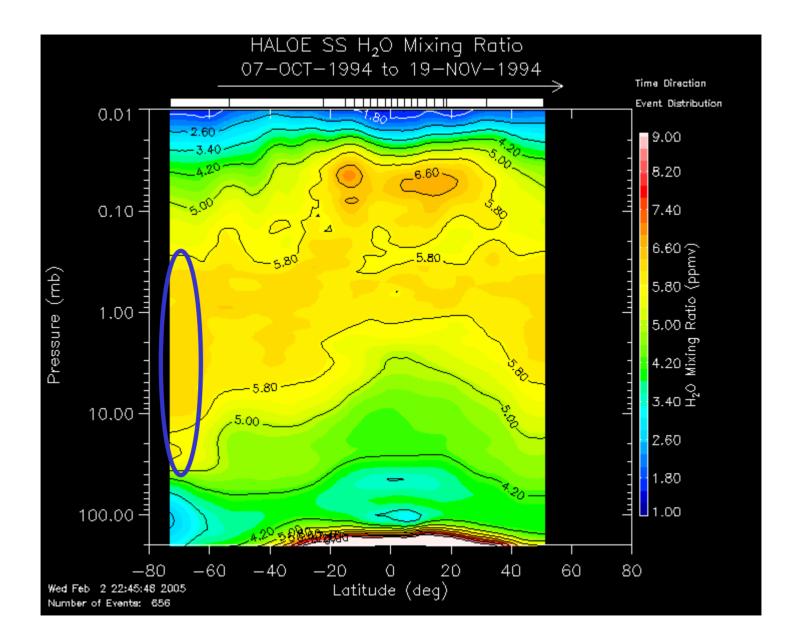
MLS 5 Day Ave. at +/- 2° of POAM Lat. POAM 5 Day Average, 3 km vertical smoothing POAM 5 Day Median, 3 km vertical smoothing

What's this?

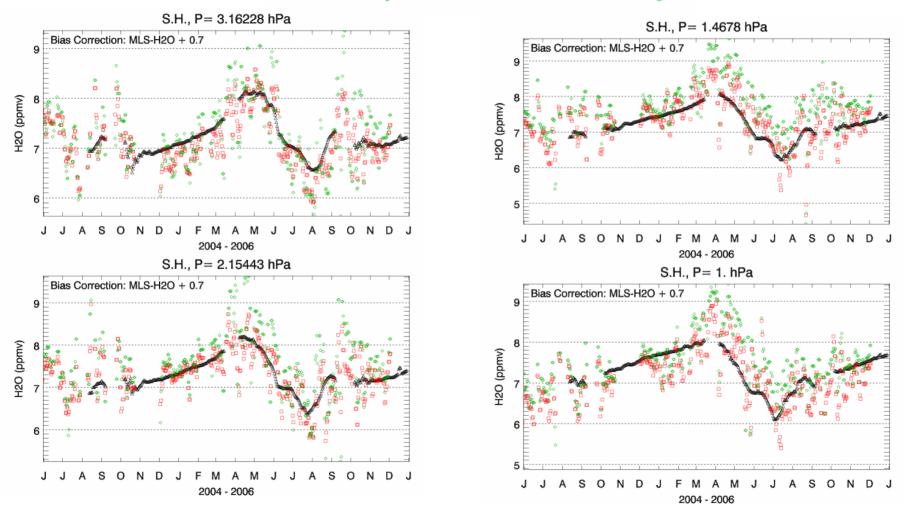


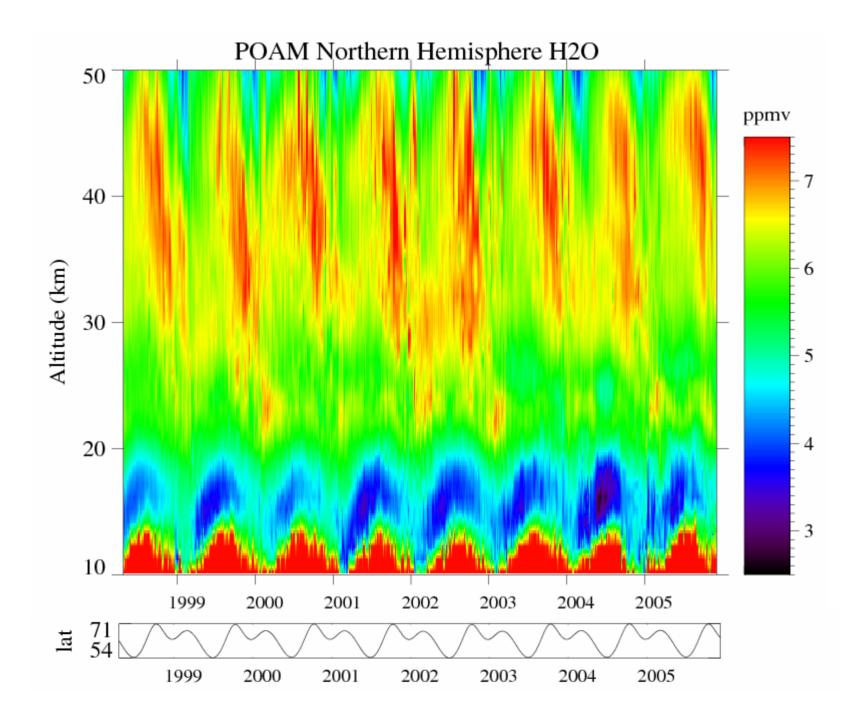




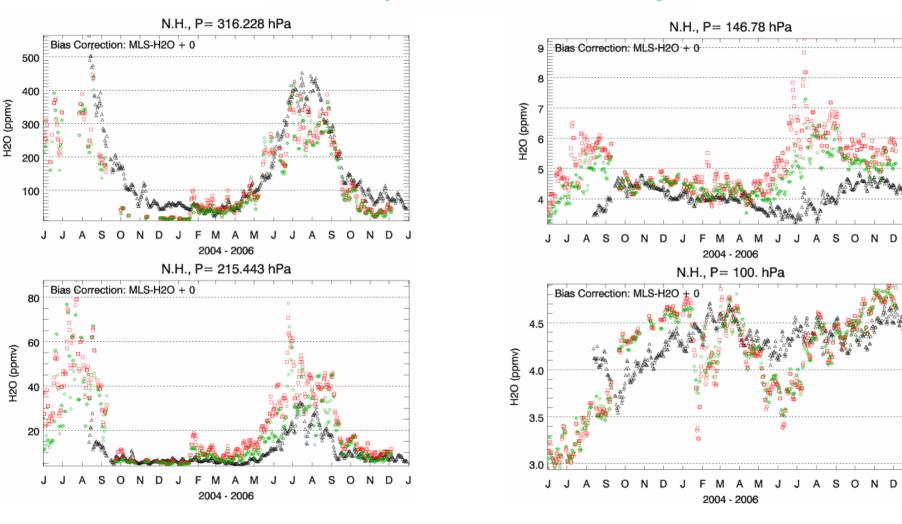


## Upper stratosphere measurements

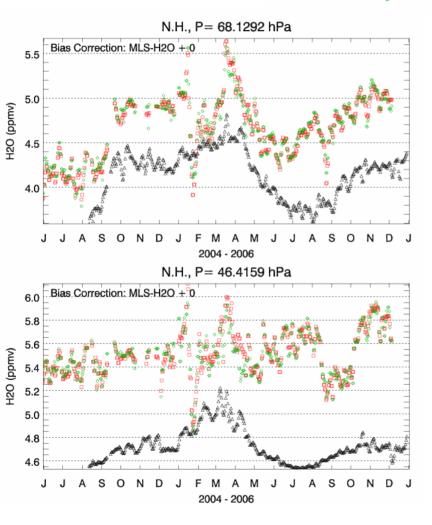


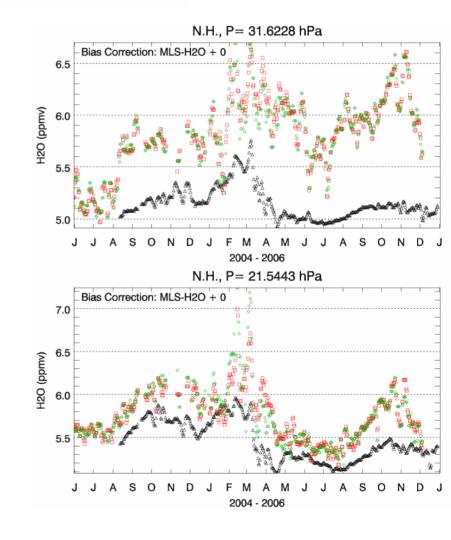


## Troposphere/lowermost stratosphere measurements

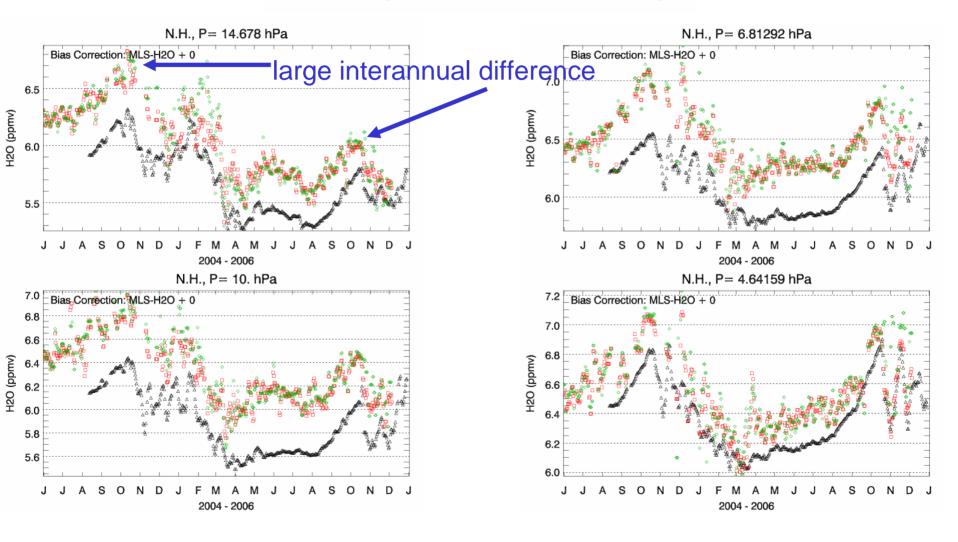


### Lower stratosphere measurements

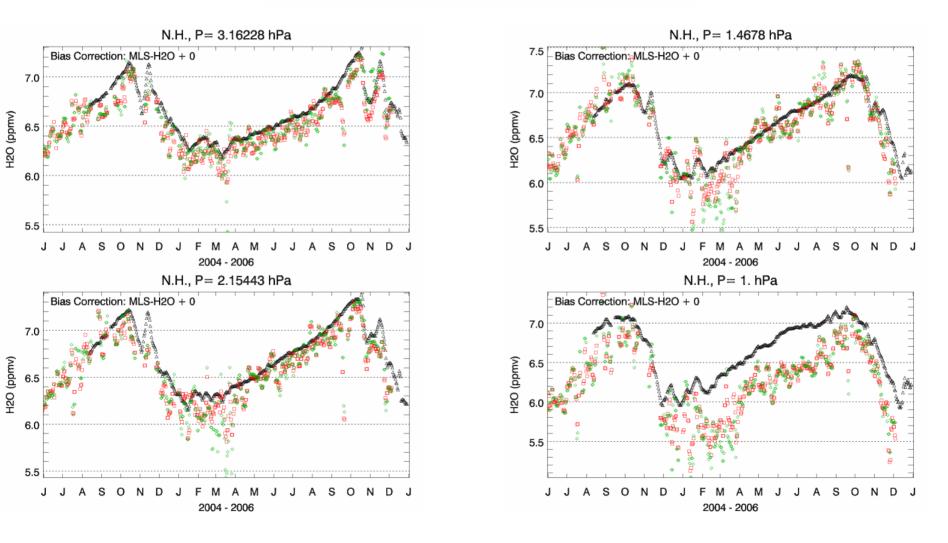




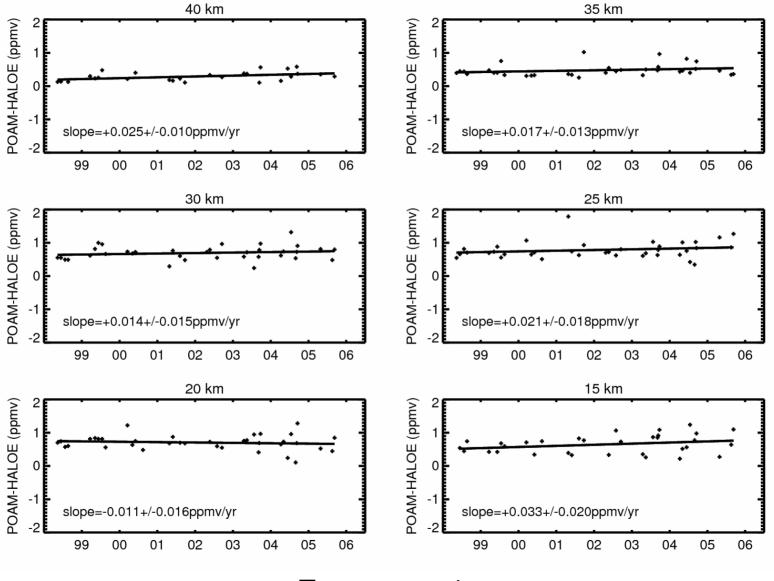
### Upper stratosphere measurements



## Upper stratosphere measurements

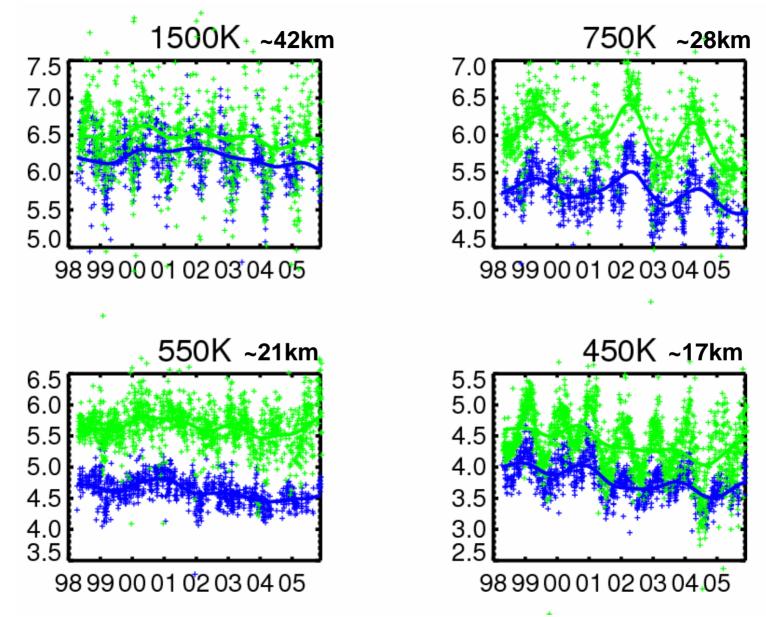


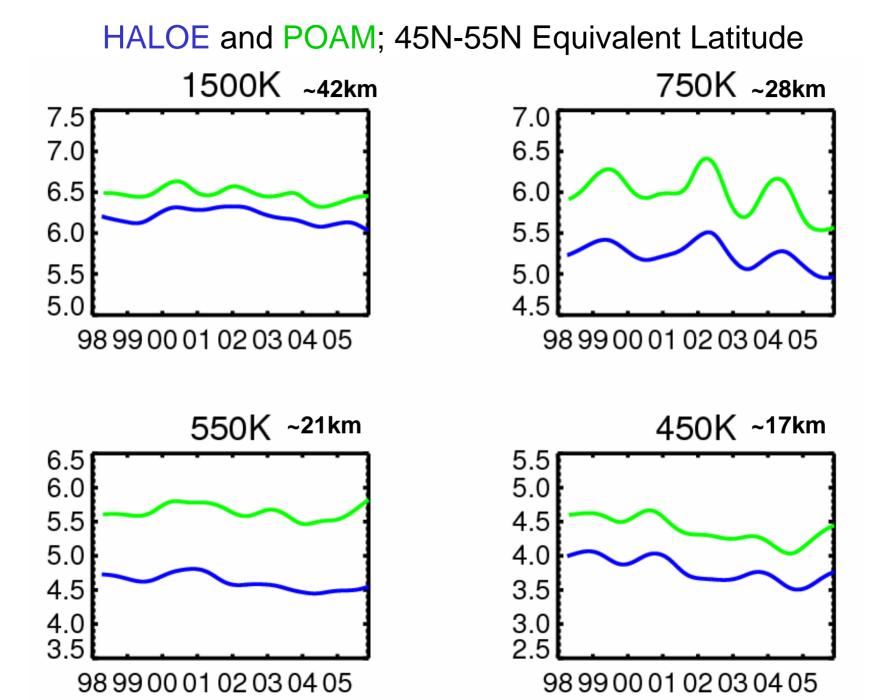
POAM sunrise - HALOE H<sub>2</sub>O comparisons 15-40km

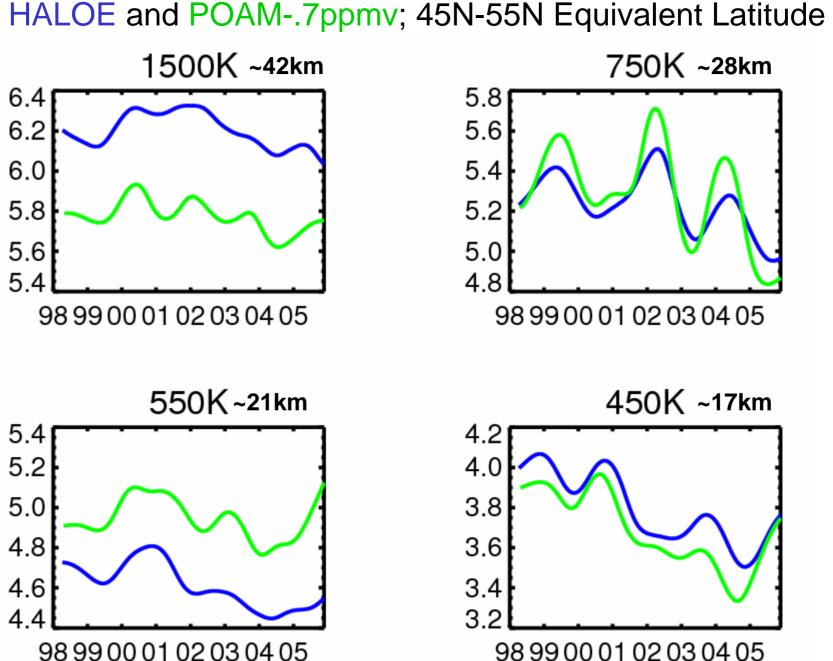


Errors are  $1\sigma$ 

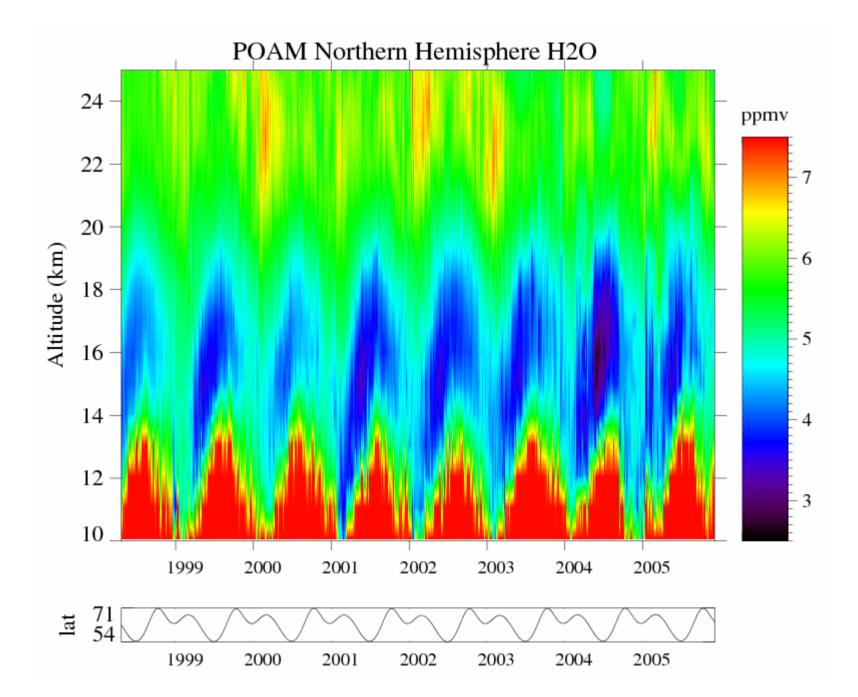






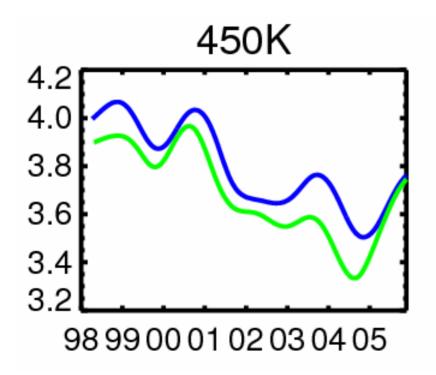


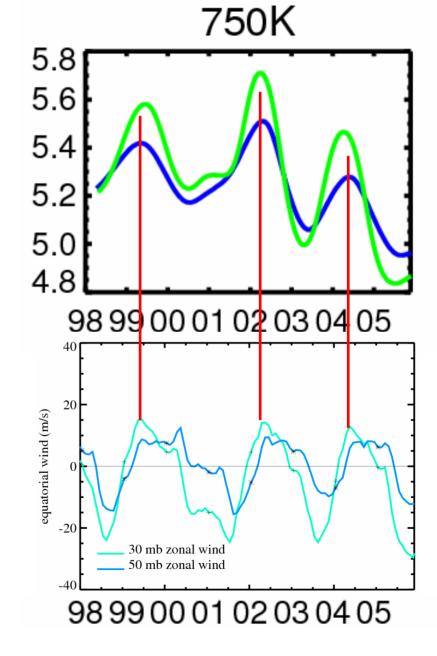
98 99 00 01 02 03 04 05



### HALOE and POAM-.7ppmv; 45N-55N Equivalent Latitude

•Good correlation between 30 and 50 mb zonal winds and 750K  $H_2O$  variations at 45-55N equivalent latitude.

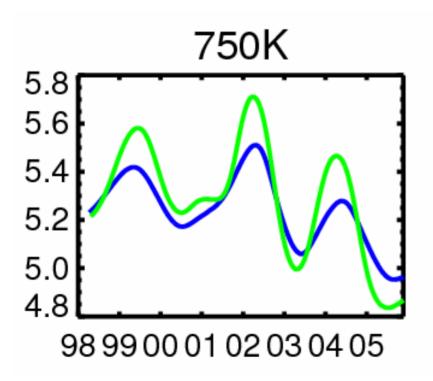


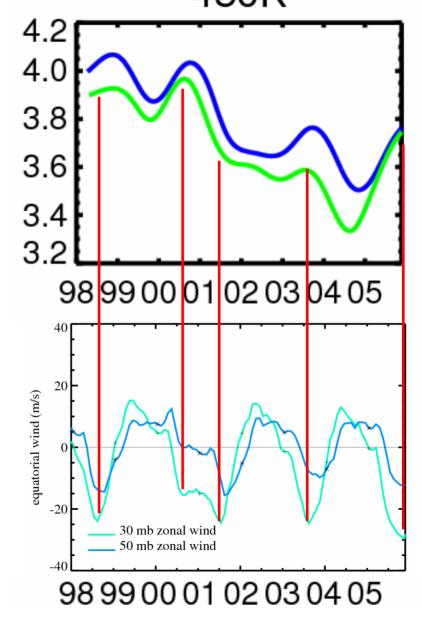


#### HALOE and POAM-.7ppmv; 45N-55N Equivalent Latitude 450K

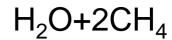
•Good correlation between 30 and 50 mb zonal winds and 750K  $H_2O$  variations at 45-55N equivalent latitude.

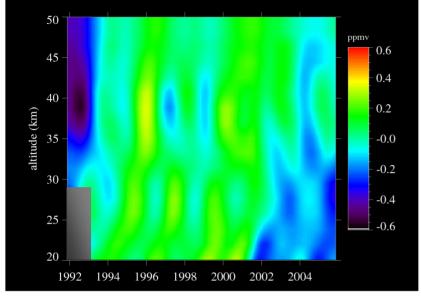
•Can the increased H2O at 450K in 2005 be completely attributed to the QBO?

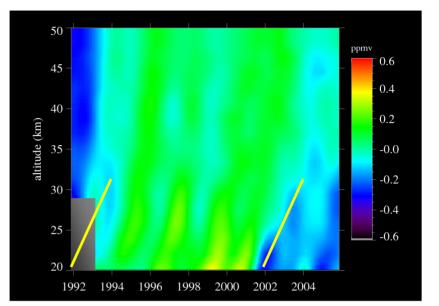




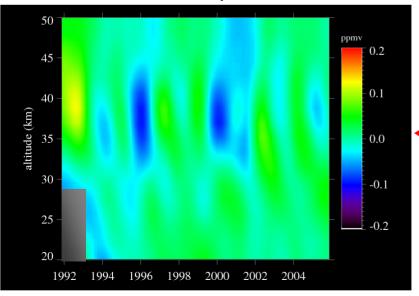
 $H_2O$ 



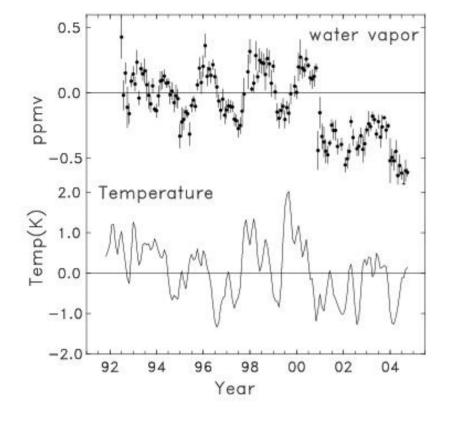




 $CH_4$ 



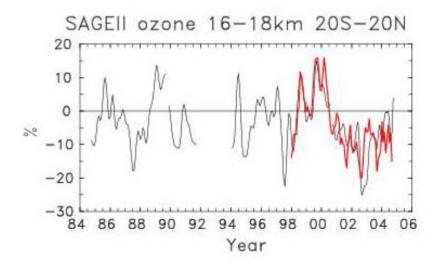
# Deseasonalized HALOE data; 20S-20N

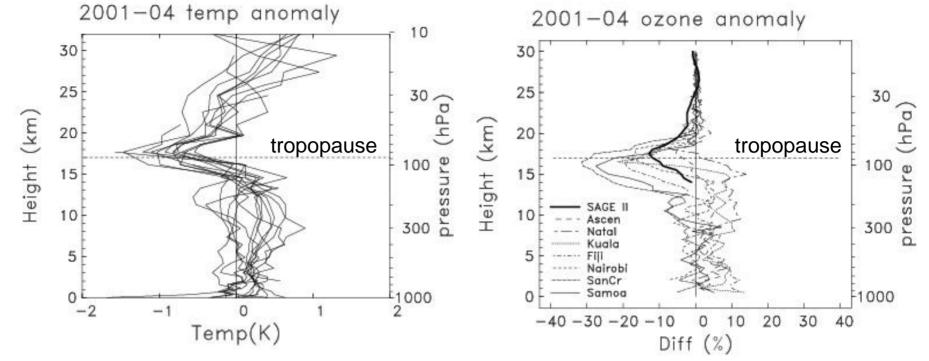


HALOE H<sub>2</sub>O (10S-10N) and deseasonalized cold point temperature anomalies from 14 radiosonde stations

(Randel et al., in press)

Zonal mean ozone anomalies 16-18km, 20S-20N. SAGE II and 7 tropical ozonesonde stations from SHADOZ





2001-2004 temperature anomaly (as compared to 1994-2000) for tropical radiosonde stations from 10S-10N

2001-2004 ozone anomaly (as compared to 1998-2000) for SAGE II and 7 tropical stations

•Decreased ozone may be the result of increased upwelling

Decreased ozone will reduce local heating

(Randel et al., in press)

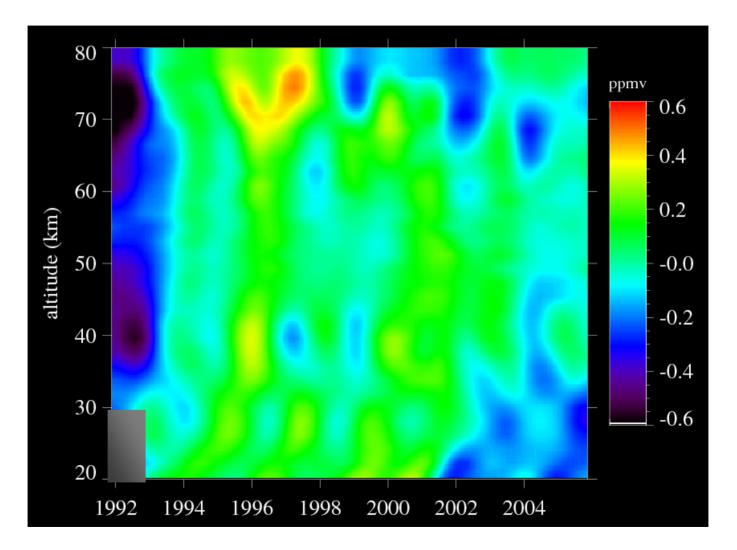
# What's next for water vapor?

- Upper stratosphere/lower mesospheric H<sub>2</sub>O constant since 1996.
- Lower stratospheric H<sub>2</sub>O down since 2001.
- Even the long-term H<sub>2</sub>O trend from CH<sub>4</sub> forcing trend is no longer clear. Will CH<sub>4</sub> continue to increase?

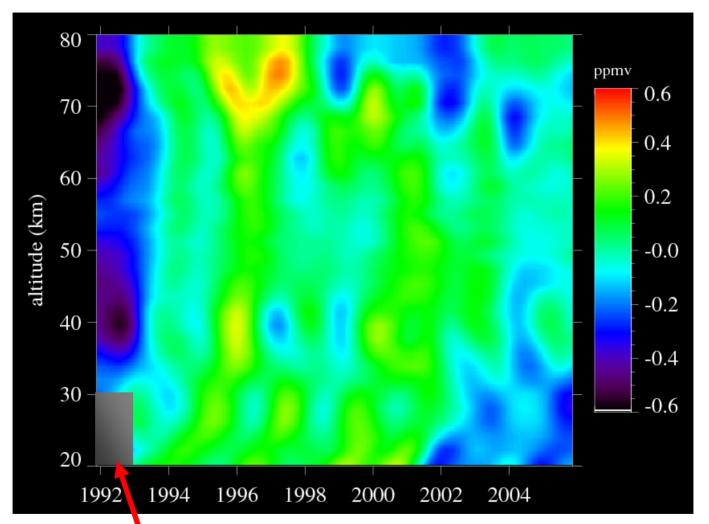
What we do and don't know about the 2001 lower stratospheric  $H_2O$  decrease and its causes

- Several instruments (HALOE, POAM III, Balloons) show decrease in water vapor.
- Decrease is correlated with decreasing tropopause temperatures.
- Ozone at the tropopause decreases at the same time as the temperature.
- Decreased ozone could be caused by increasing tropical upwelling, but there is no other evidence of changes in circulation.

### Deseasonalized HALOE water vapor; 20S-20N



### Deseasonalized HALOE water vapor; 20S-20N



Lower stratospheric water vapor in 1991-1992 probably affected by Pinatubo aerosols. Unusually dry compared to MLS (SPARC report).

## Water vapor fits

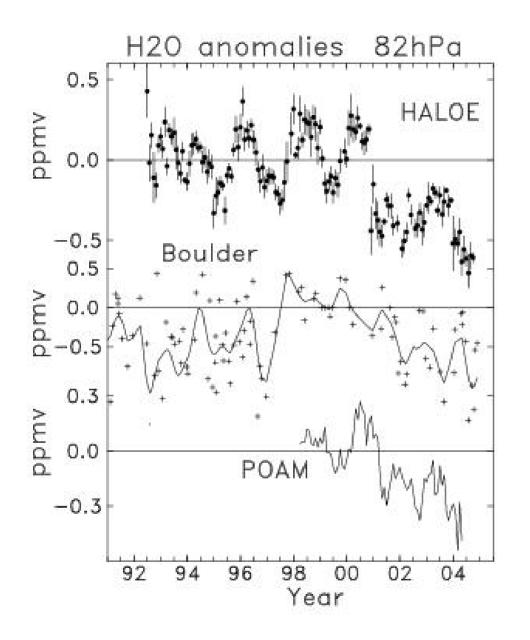
## Fit using (t in years):

 $f = a_0 +$ (annual)  $a_1 \sin(2\pi t) + a_2 \cos(2\pi t) + a_2 \cos(2\pi t)$ (semi-annual)  $a_{3}\sin(4\pi t) + a_{4}\cos(4\pi t) + a_{4}\cos(4\pi t)$  $a_5 \sin((12/27)2\pi t) + a_6 \sin((12/27)2\pi t)$ (QBO) (linear trend)  $+ a_7 t$ Or (solar cycle term) +a<sub>7</sub>F(Mg II)

# Comparing the early 1990's increase with the 2001 decrease

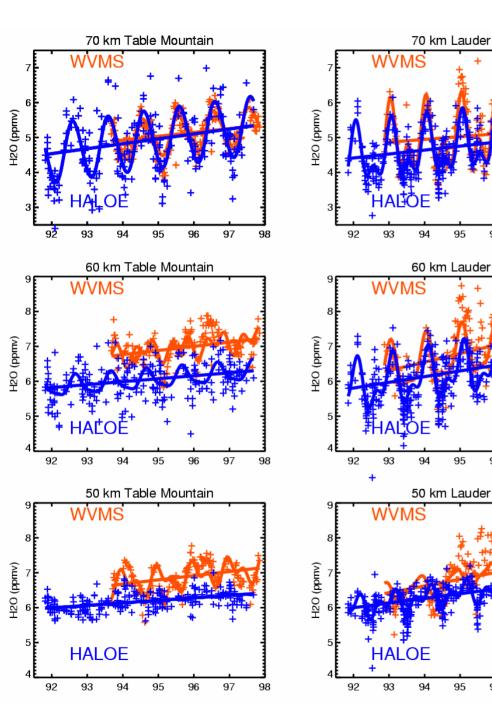
- Early 1990's increase
  - Water in lower mesosphere increased by ~0.5 ppmv over 5 years
  - No reliable measurements in the lower stratosphere available
  - Measured tropopause temperatures not correlated with increase
  - Water vapor remained at new level until 2001

- 2001 decrease
  - Slow ascent of signal:
     from ~20km to ~30km in
     2 years
  - Water vapor in the lower stratosphere shows a decrease of ~0.4 ppmv
  - Measured tropopause temperatures well correlated with decrease
  - Occurred after a long QBO cycle
  - Water vapor remains at new level in 2005 (2006?)



Balloon data shows a similar decline to HALOE and POAM

(Randel et al., in preparation)



Early 1990's trend and seasonal cycle fit WVMS H<sub>2</sub>O trends for 40-60km (1992-1997) ~0.15 ppmv/year Global HALOE H<sub>2</sub>O tre

96

96

96

97

98

97

98

97

Global HALOE H<sub>2</sub>O trend for 40km-60km (1991-1997) ~0.13 ppmv/year

Of this, ~0.04 ppmv/year is attributable to changes in  $CH_4$  conversion to  $H_2O$ 

# Some final thoughts

- While there where good reasons for expecting an increase in H<sub>2</sub>O following the eruption of Pinatubo because of aerosol loading and the resultant warming of the tropopause, we don't understand why the H<sub>2</sub>O mixing ratio remained high.
- The H<sub>2</sub>O decrease in 2001 is consistent with tropopause temperature changes, but we don't understand why this temperature changed and then remained low.
- CH<sub>4</sub> is likely to increase over time, so this will contribute to a long-term increase in stratospheric H<sub>2</sub>O, but this trend is likely to be very slow and will, especially in the lower stratosphere where only a fraction of the CH<sub>4</sub> has been oxidized. On decadal timescales this trend will probably be overwhelmed by other changes.

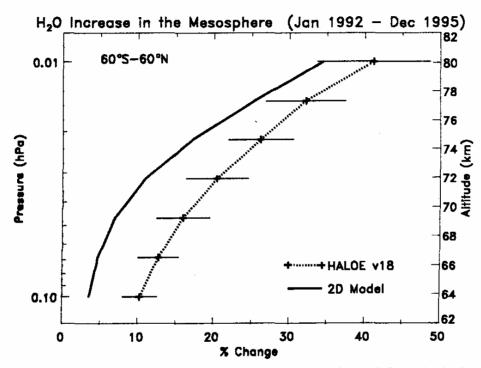
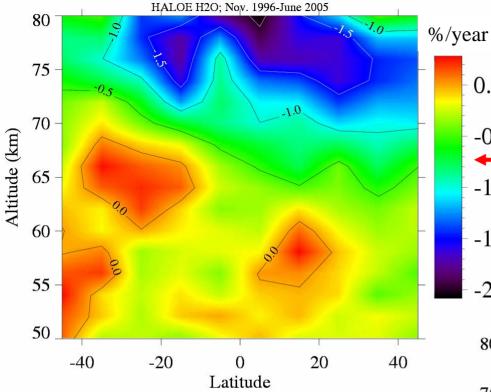


Figure 4. Height profiles of the 4 year trend estimated from both the model and the HALOE data averaged over  $60^{\circ}$ S to  $60^{\circ}$ N.

From Chandra et al., 1997



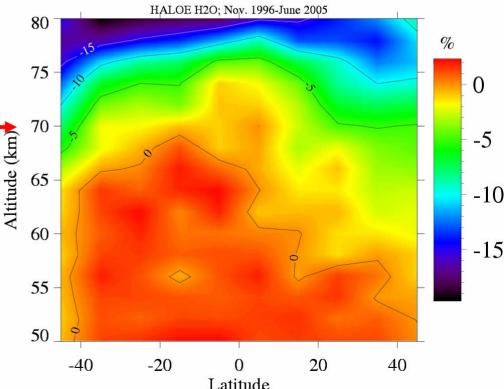
Now replace the trend fit term • with a fit to a normalized Mg II index.

-15% =>  $\sim$ 30% peak-to-peak change in H<sub>2</sub>O from solar min to solar max. har HALOE H<sub>2</sub>O fits binned by latitude
 0.0 Nov. 1996-June 2005

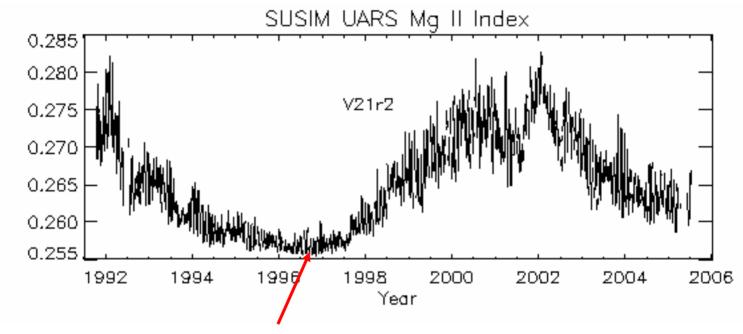
-1.5

-2.0

-0.5
%/year trend calculated without a
-1.0 Mg II index term included in fit



# 11 year Solar Cycle



### minimum Mg II in 1996 => minimum Lyman-α photodissociation => water vapor maximum in mesosphere

http://wwwsolar.nrl.navy.mil/susim\_uars\_mgii\_index.html

#### Deseasonalized HALOE H<sub>2</sub>O and Mg II fit

