



# Water Vapor Mm-wave Spectrometer

Gerald E. Nedoluha, R. Michael Gomez,  
Brian C. Hicks, Richard M. Bevilacqua  
*Naval Research Laboratory*





## Water Vapor Mm-wave Spectrometer (WVMS)



22 GHz radiometers

Water vapor profile measurements from 40-80 km

Measurements from 3 NDSC sites:

Lauder, New Zealand ( $45^{\circ}\text{S}$ ,  $169.7^{\circ}\text{E}$ ):

Nov. 1992-Apr. 1993, Jan. 1994-present

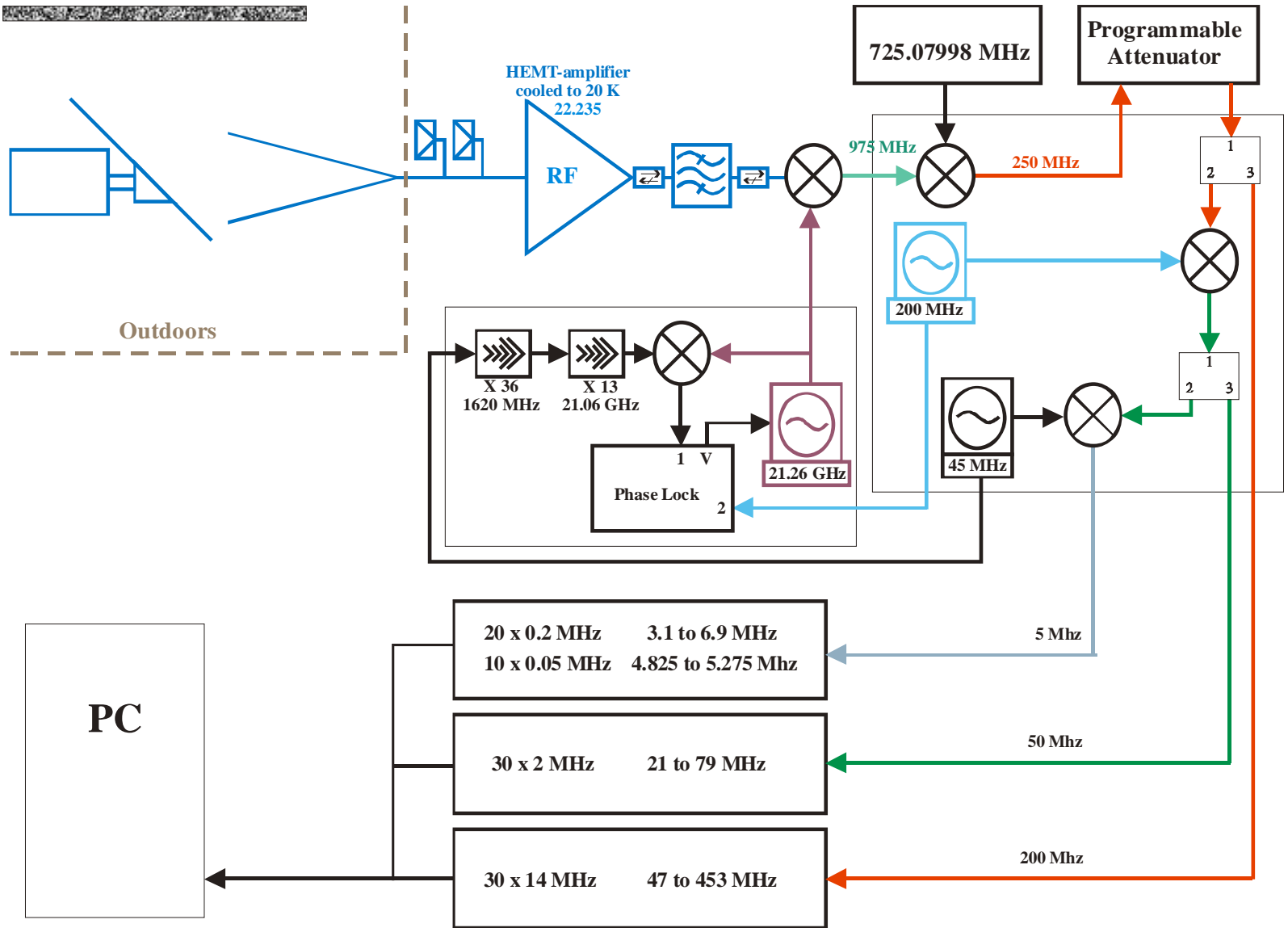
Table Mountain, CA ( $34.4^{\circ}\text{N}$ ,  $242.3^{\circ}\text{E}$ ): May 1993-Nov. 1997

Mauna Loa, HA ( $19.5^{\circ}\text{N}$ ,  $204.4^{\circ}\text{E}$ ): Mar. 1996-present



# Instrument description

- Front end
  - Corrugated feed horn with beam size  $\sim 8^\circ$  FWHM
  - Rotating aluminum plate mirror at  $45^\circ$  to feedhorn
  - Fixed absorber bar used for noise balance
  
- Back end
  - HEMT amplifier (20K)
  - Noise diode calibration source
  - Spectrometer centered at 22.235 GHz
  - Single side-band
  - New spectrometers (WVMS3 Mauna Loa, WVMS2 Table Mountain)
    - 10x50kHz, 20x200kHz, 30x1MHz, 30x14MHz
  - Old spectrometer (WVMS1 Lauder)
    - 20x200KHz, 20x2MHz, 10x40MHz

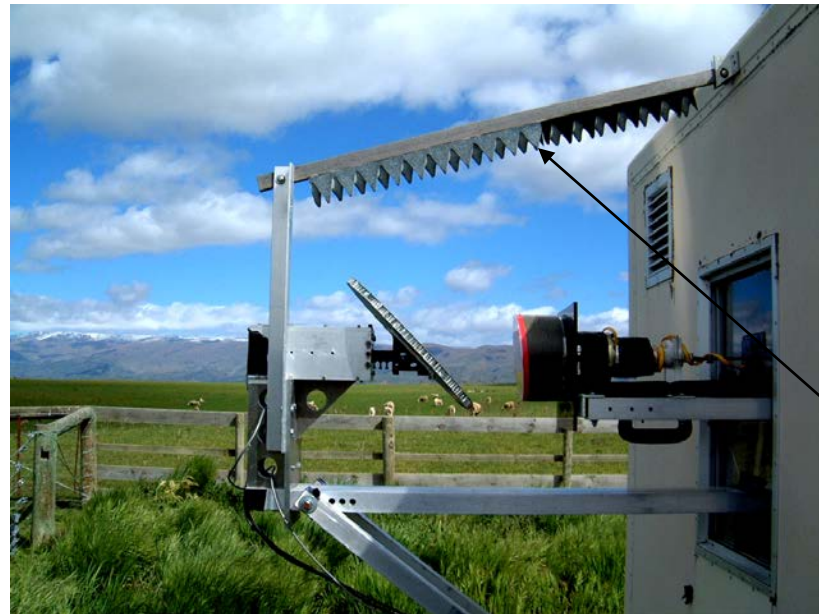




# Calibrating the noise diode



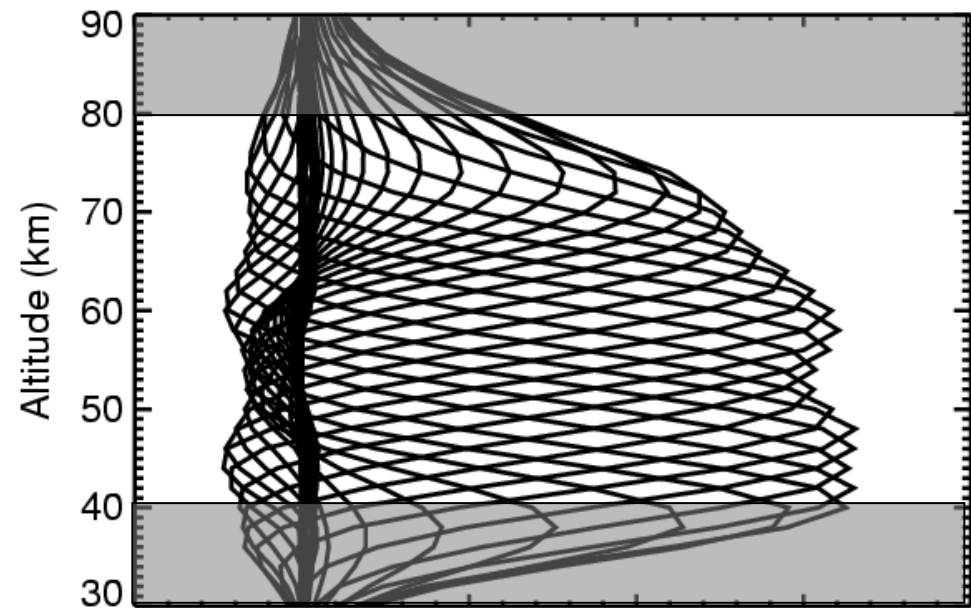
Mirror and feedhorn close-up



Bar for noise balancing

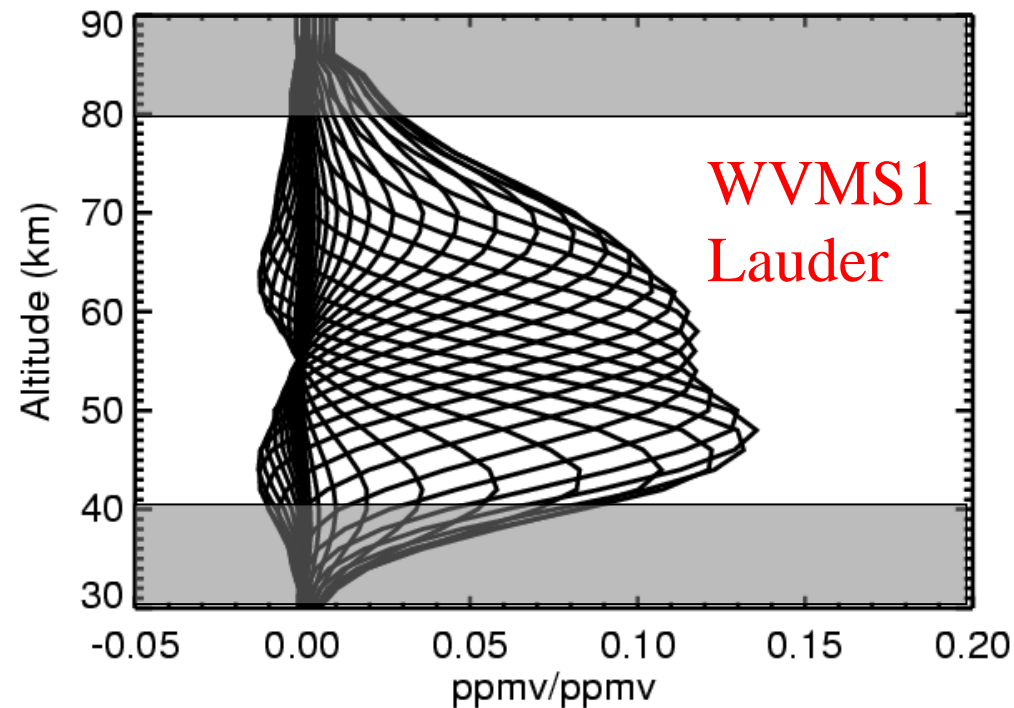
# Data taking procedure

- Tipping measurement
  - Using angles from 45° (limited by absorber bar) to 75° (limited by horizon)
  - Find least-squares fit to
$$T_{\text{sys}} = T_{\text{rx}} + T_{\text{atm}}[1 - \exp(-\mu\tau_{\text{trop}})]$$
solving for  $T_{\text{sys}}$ ,  $T_{\text{rx}}$ , and  $\tau_{\text{trop}}$
- Find a signal angle such that  $T_{\text{sig}} \sim T_{\text{ref}}$ , where  $T_{\text{ref}}$  is measured near zenith and includes a contribution from the absorber bar
- Make 5 sets of 3 measurements and average together
  - $T_{\text{sig}}$  with noise diode off
  - $T_{\text{ref}}$  with noise diode off
  - $T_{\text{ref}}$  with noise diode on



WVMS3  
Mauna Loa

Averaging kernels for 500 scan integrations from WVMS3 at Mauna Loa, and WVMS1 at Lauder.



WVMS1  
Lauder

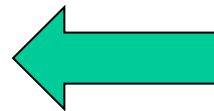
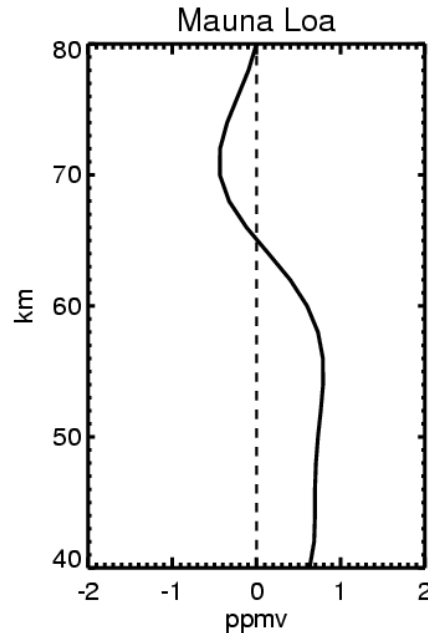
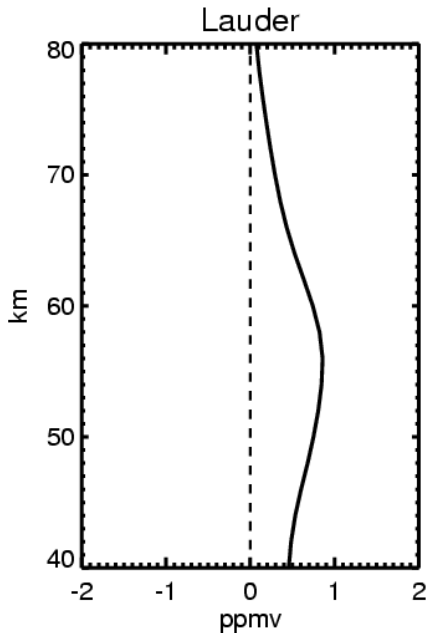
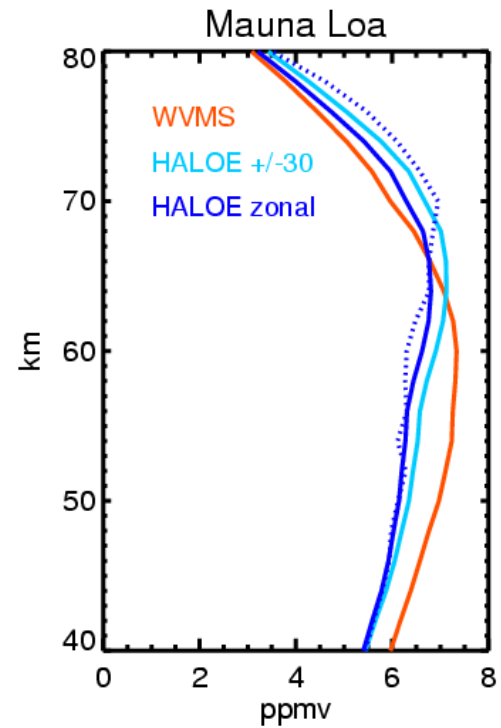
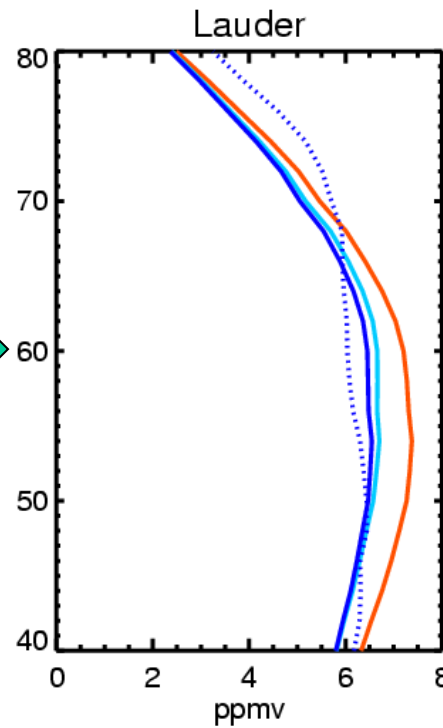
Note the difference in the sensitivity of the 2 instruments

All HALOE data compared with WVMS is convolved with these averaging kernels



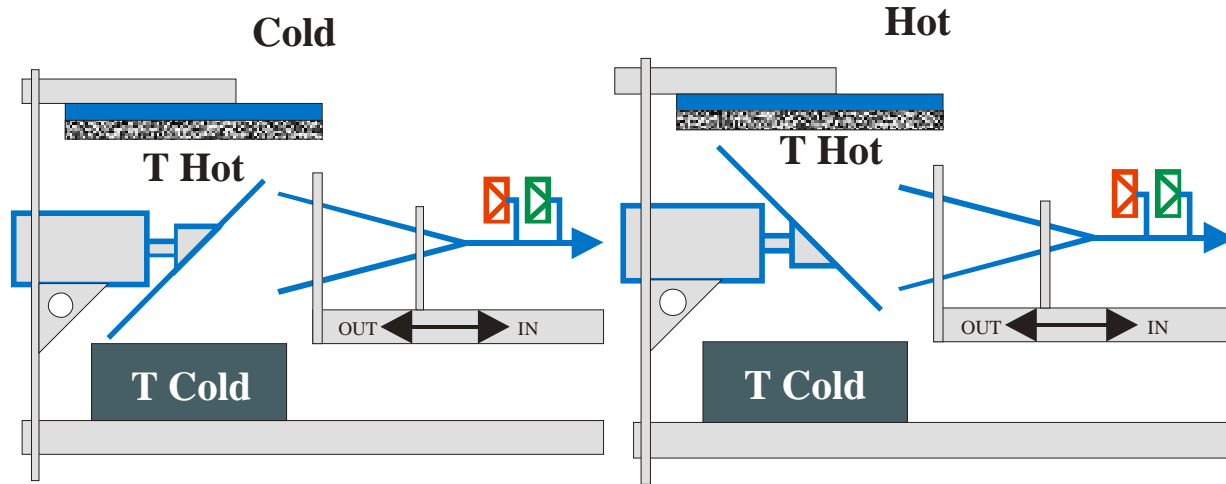
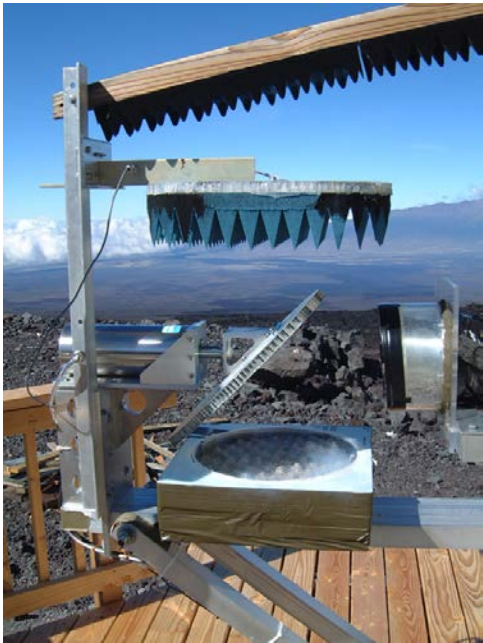
Average of coincident measurements:

- **WVMS**
- **HALOE** (+/- 5° latitude +/- 30° longitude)
- **HALOE zonal** (+/- 5° latitude)
- **unconvolved HALOE zonal** (dots)



**WVMS-HALOE**





$$\text{Cold 1} = (\text{out} + \text{in}) / 2$$

$$\text{Cold ND1} = (\text{out} + \text{in}) / 2$$

$$\text{Cold ND2} = (\text{out} + \text{in}) / 2$$

$$\text{Cold 2} = (\text{out} + \text{in}) / 2$$

$$\text{Hot} = (\text{out} + \text{in}) / 2$$

$$\text{Noise Diode 1 Temperature} = \frac{(\text{Cold ND1} - \text{Cold1}) * (\text{T Hot} - \text{T Cold})}{(\text{Hot} - \text{Cold 2})}$$

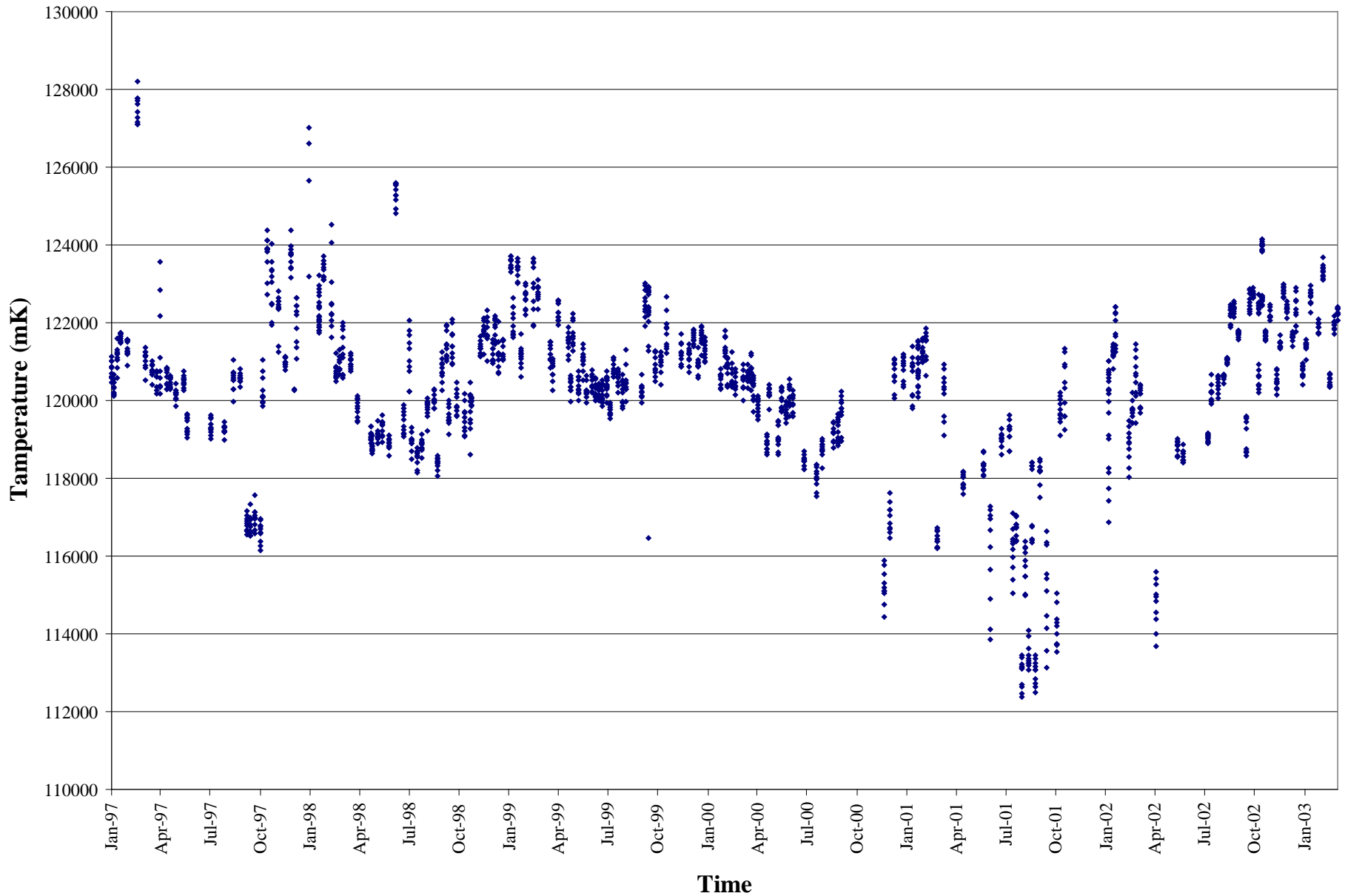
$$\text{Noise Diode 2 Temperature} = \frac{(\text{Cold ND2} - \text{Cold1}) * (\text{T Hot} - \text{T Cold})}{(\text{Hot} - \text{Cold 2})}$$

$$\text{Receiver Temperature} = \frac{(\text{Cold1}) * (\text{T Hot} - \text{T Cold})}{(\text{Hot} - \text{Cold 2})} - \text{TCold}$$

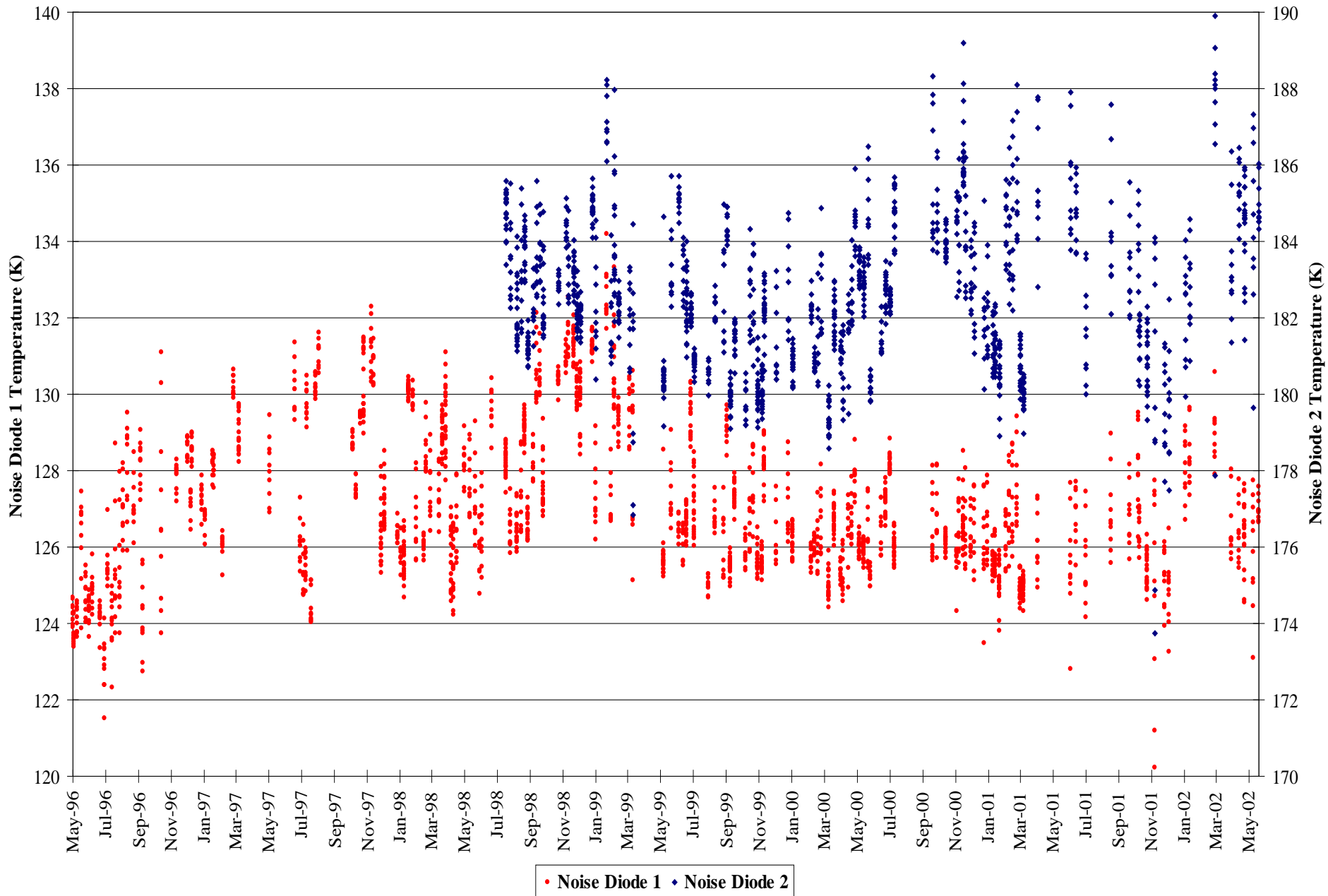
# Calibration

- Calibrate the noise diode ~weekly using a liquid nitrogen load
  - Assume that variations between weekly calibration are mainly NOT caused by real variations in the noise diode signal, but watch for change over prolonged periods
- All measurements rely on the stability of the noise diode

# WVMS1 Noise diode calibration



# WVMS3 Noise diode calibrations



# What can cause a bad calibration?

- Peaks of cold load above liquid nitrogen (windy)
- Hot load temperature not correctly measured (non-uniform temperature)
  - Brought outside from hot container and not allowed to equilibrate
  - Direct sun



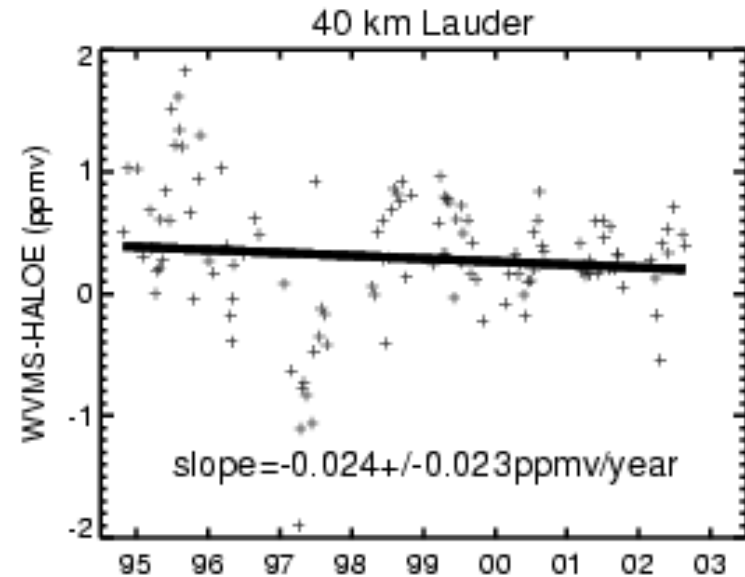
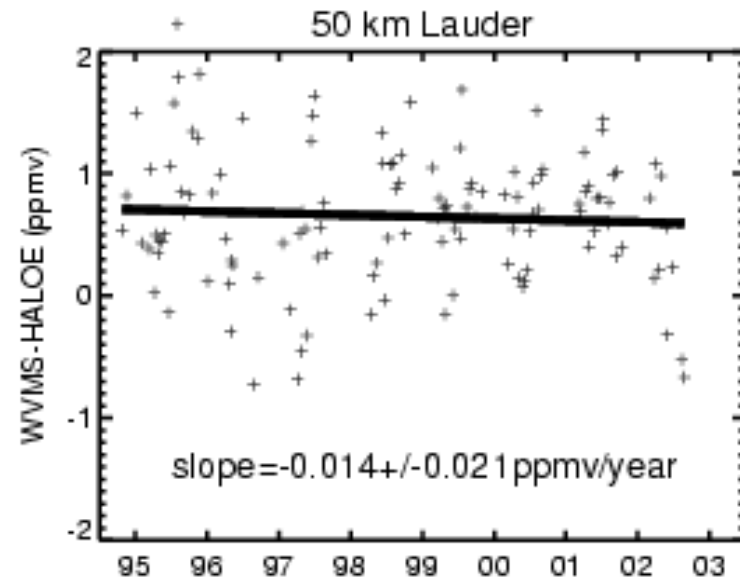
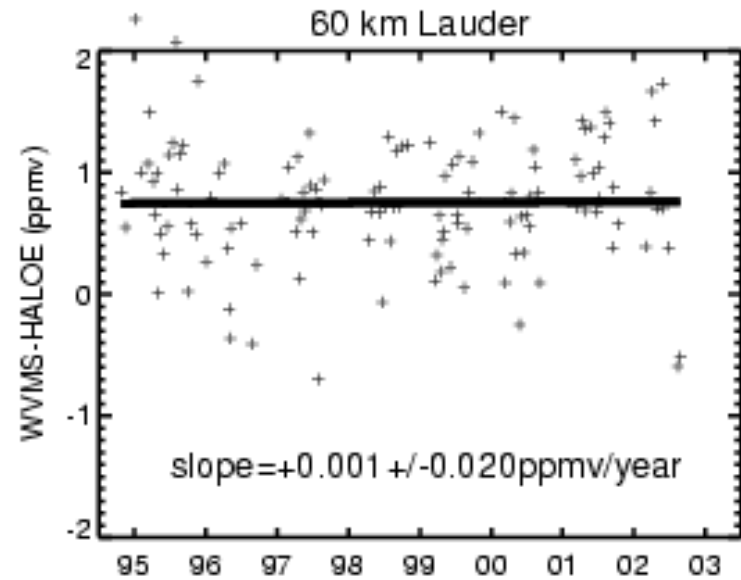
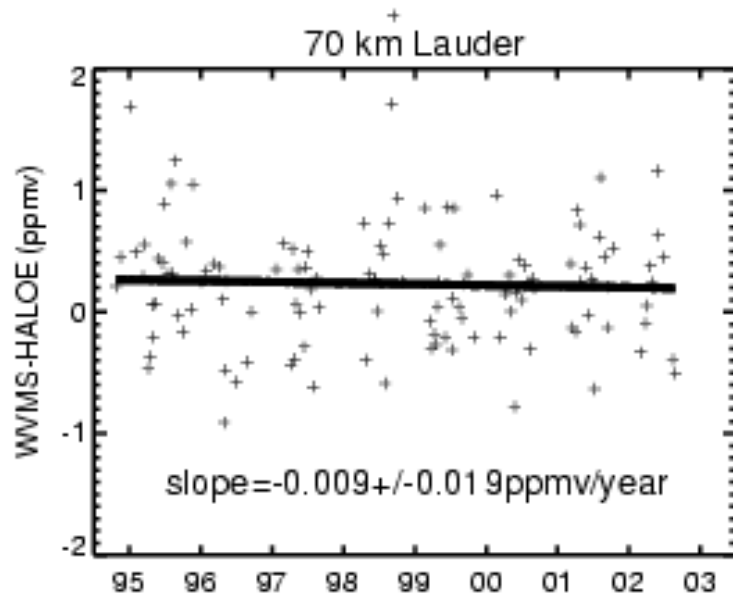
# Trend detection with the WVMS instruments

- Calibration
  - Individual measurements are calibrated against a noise diode
  - Noise diode is calibrated weekly using liquid nitrogen
- Pointing
  - 1° error in pointing causes an error ~5% in retrieved water vapor
  - Laser alignment technique used to ensure accurate pointing
- Consistency
  - Improvements to stability are always good
  - Minimal changes made to sensitivity

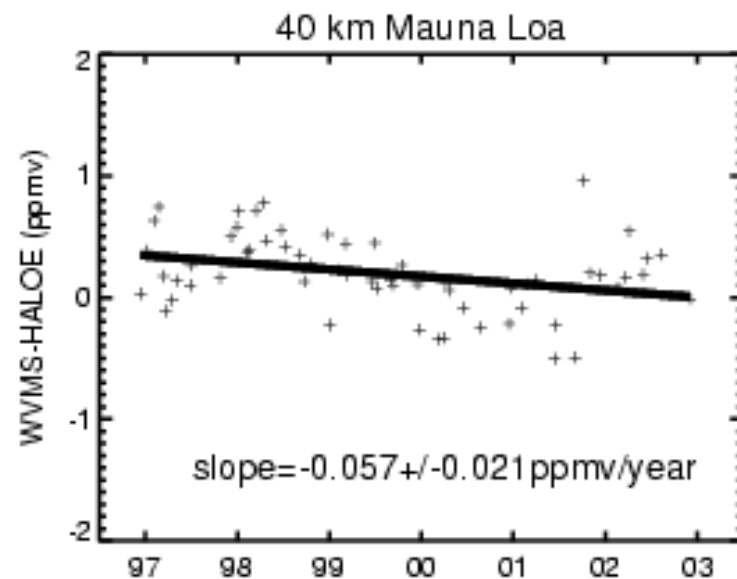
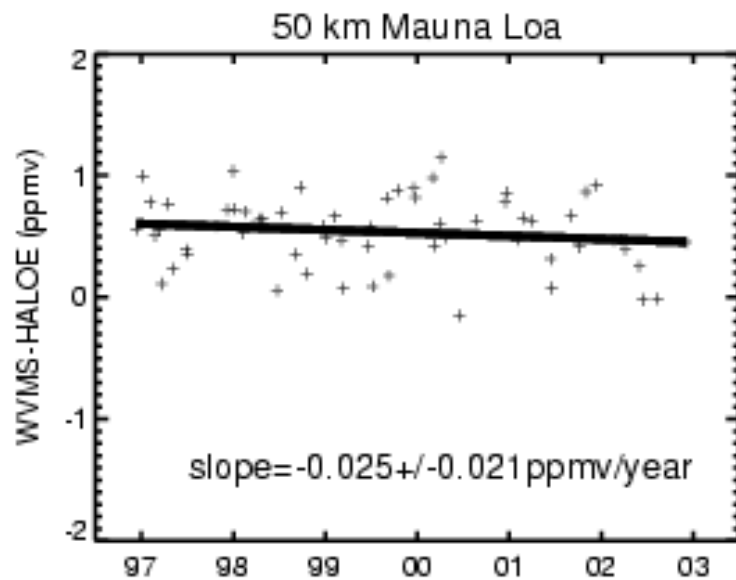
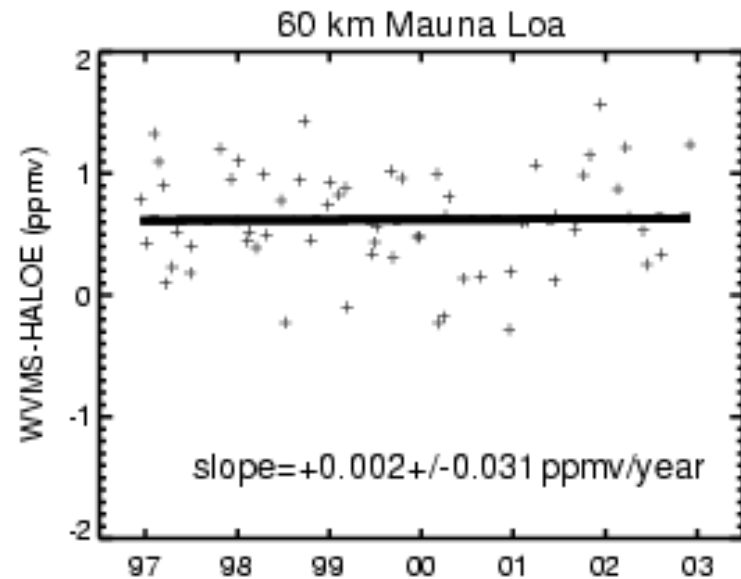
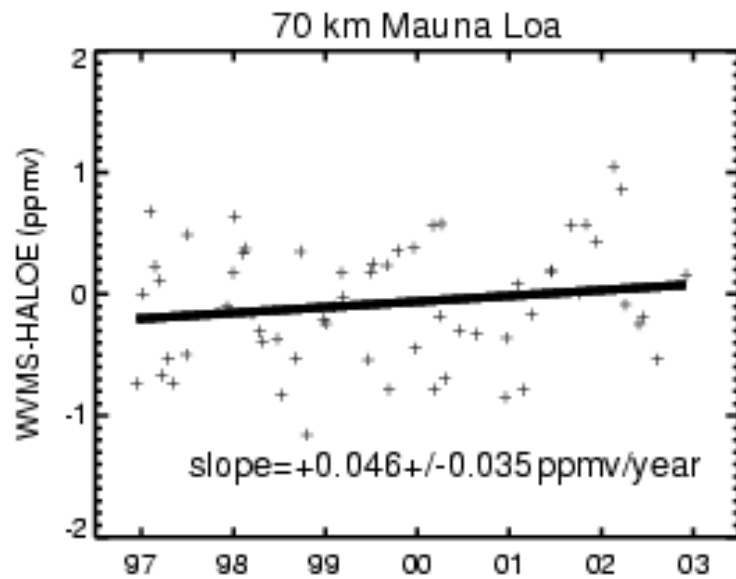
# Hard choices

- Vary signal angle or reference angle?
  - Varying the reference angle is intuitively better because this allows one to optimize the signal angle and hence maximize  $T_{\text{sig}}$ , but baselines tend to be reference angle dependent.
- How close to the horizon should we look?
  - Lower angles generally give more signal, but measurement becomes more sensitive to errors in pointing and to uncertainty in  $\tau_{\text{trop}}$

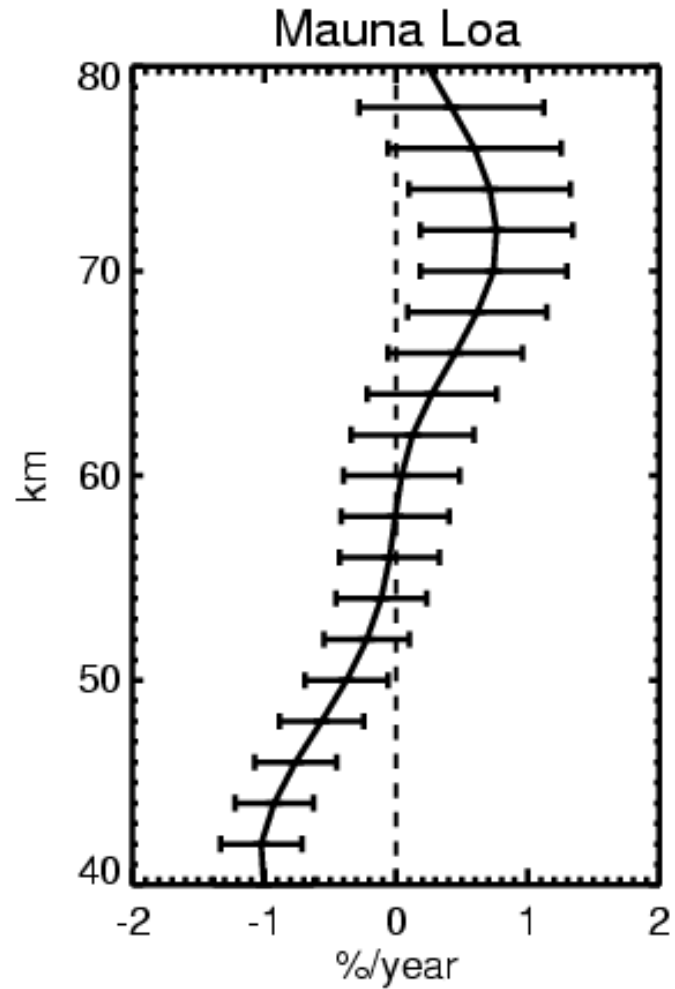
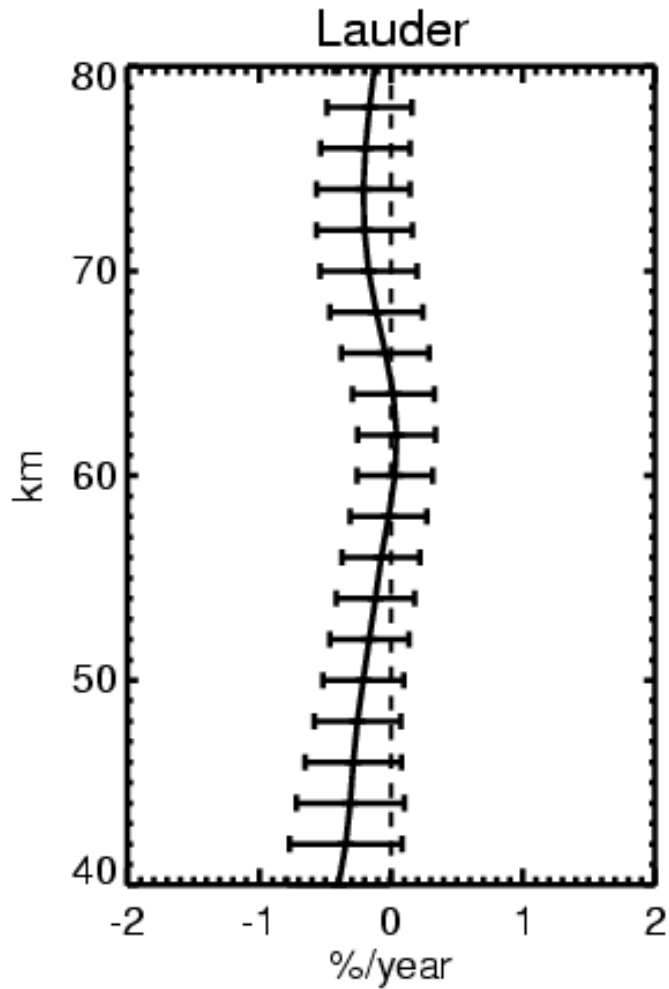
# Coincident WVMS-HALOE comparisons at Lauder, and best-fit slopes



# Coincident WVMS-HALOE comparisons at Mauna Loa

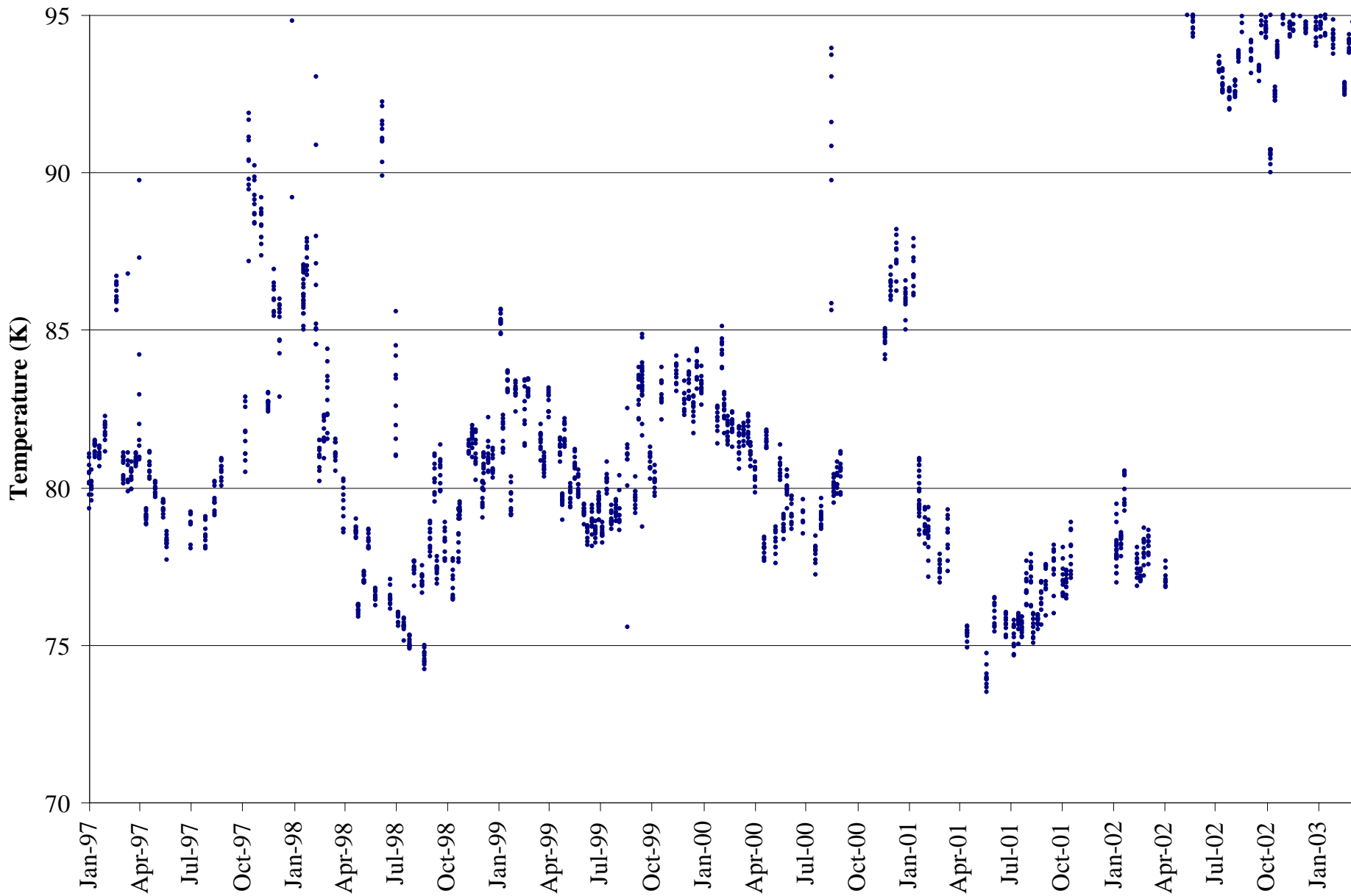


# WVMS-HALOE trends from coincident measurements

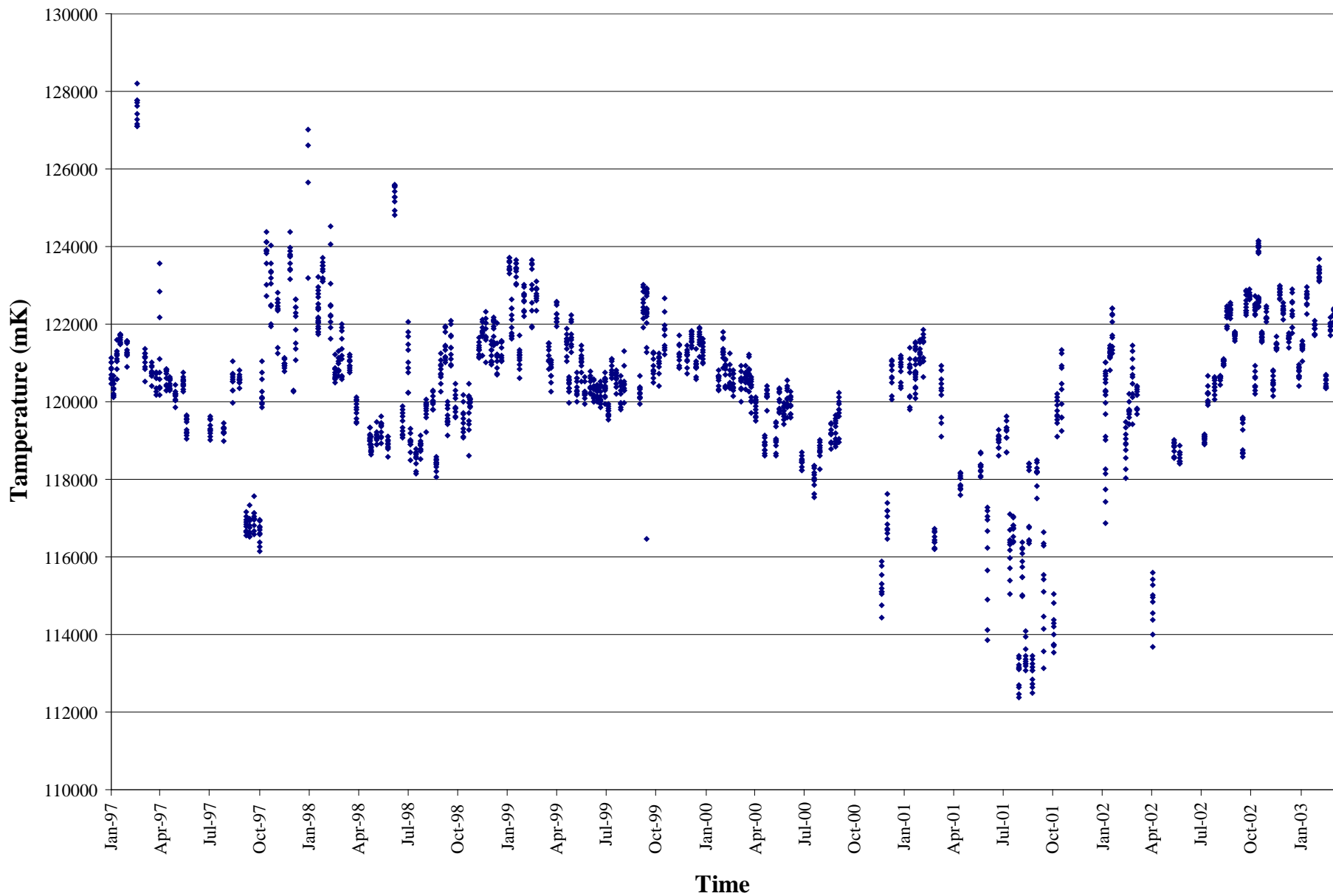




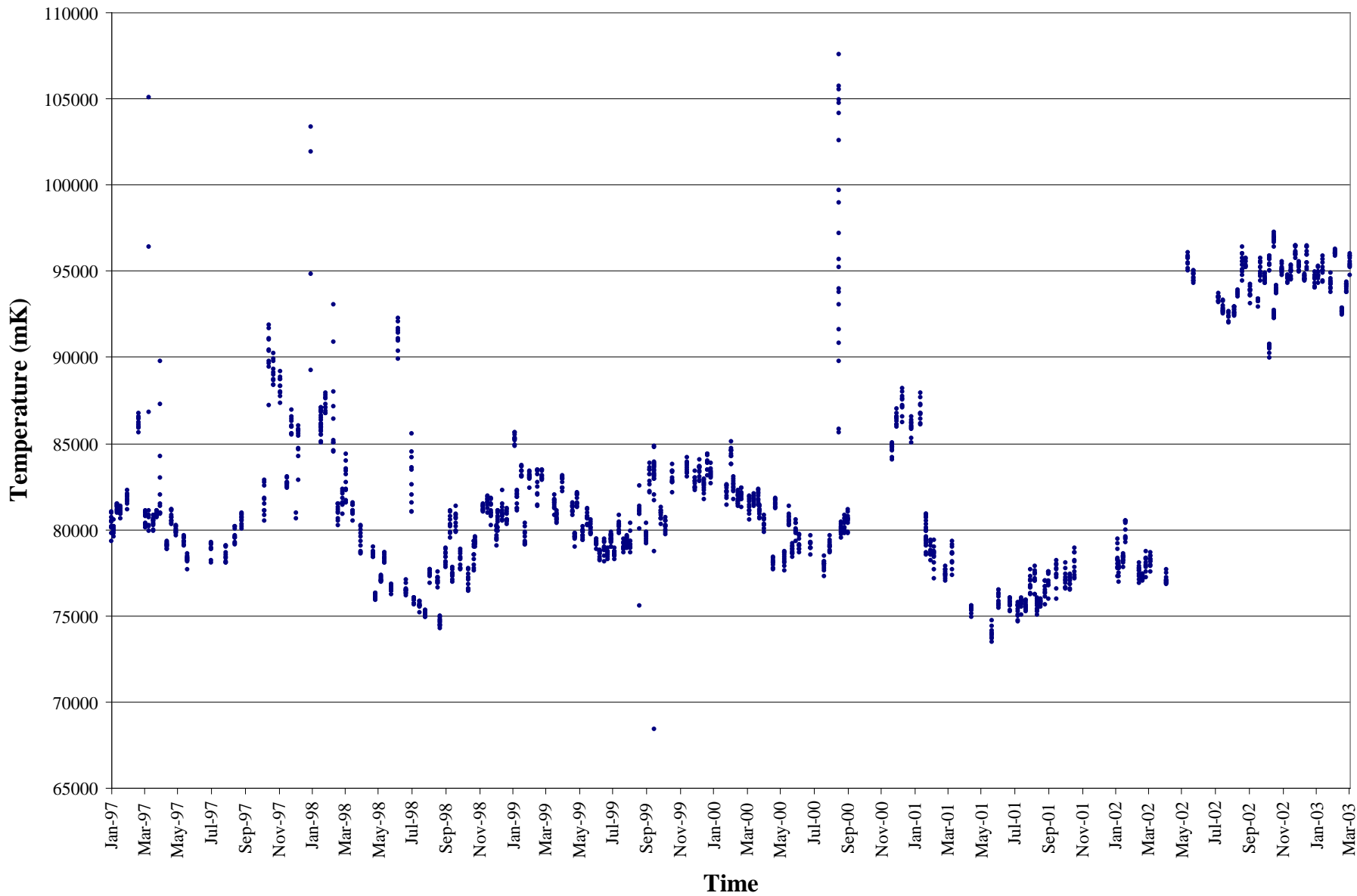
# WMS1 Receiver Temperature from Calibrations Lauder New Zealand



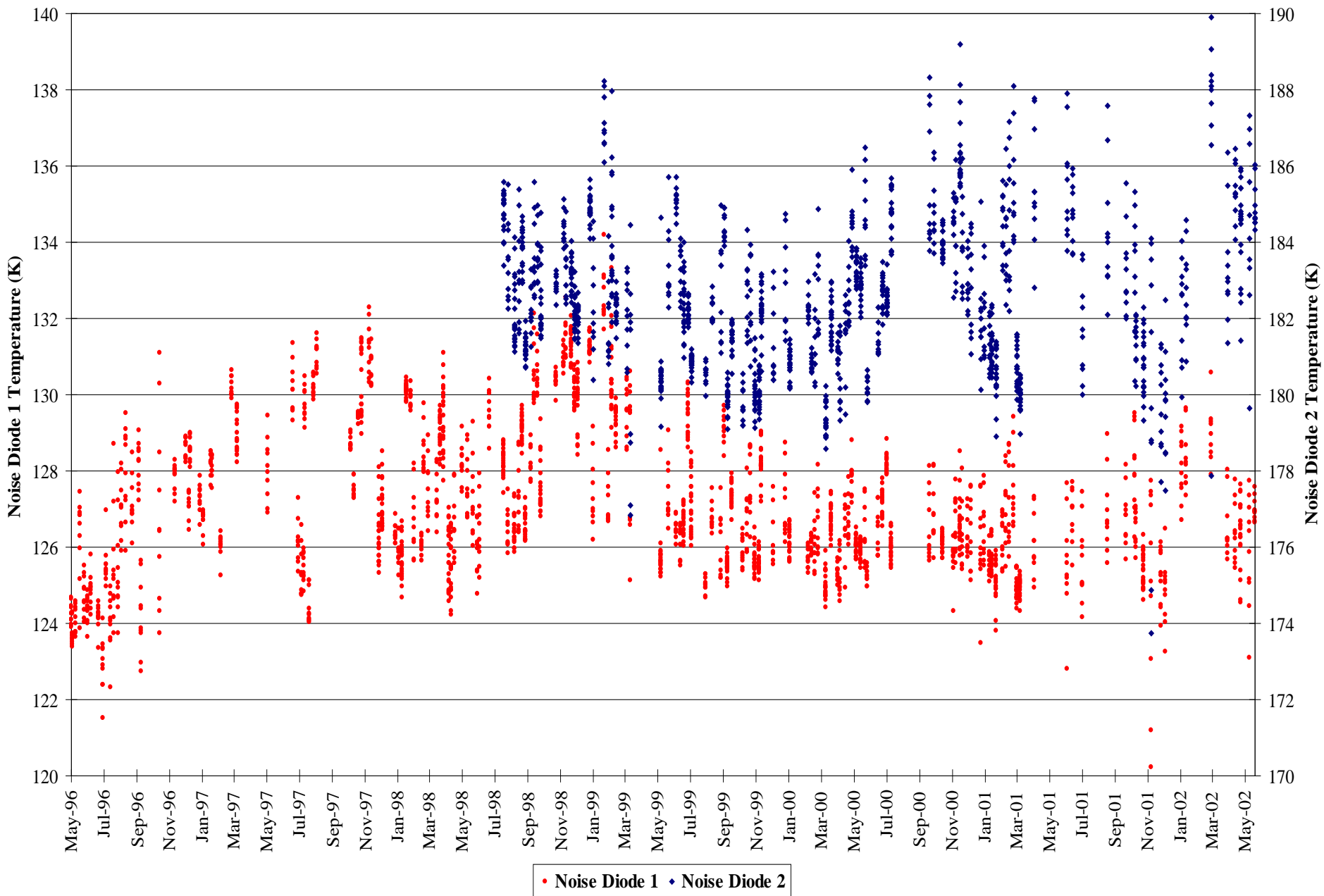
# WVMS1 Calibration Data



# WVMS1 Receiver Temperature from Calibrations

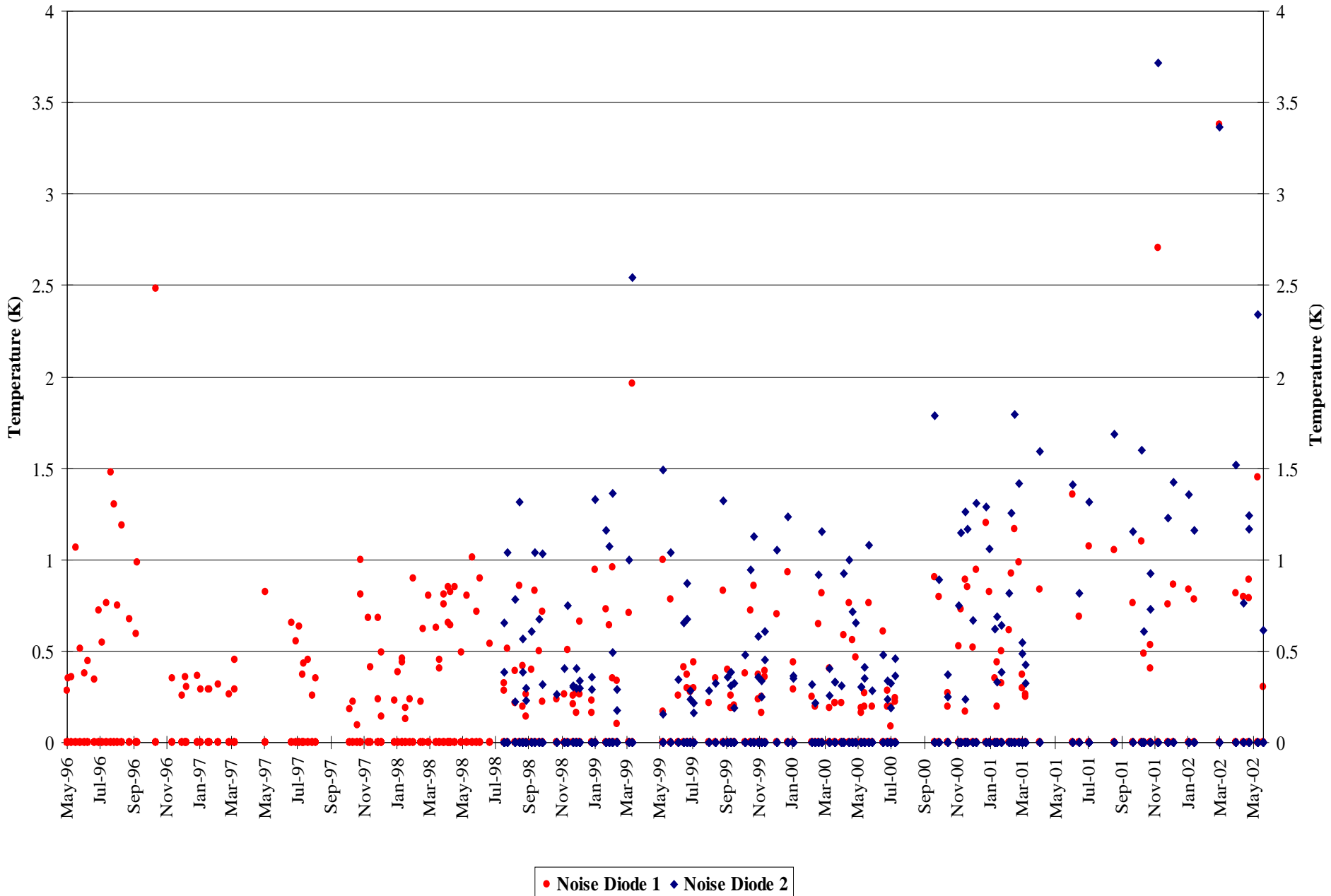


# WVMS3 Noise Diode Calibrations Mauna Loa Hawaii



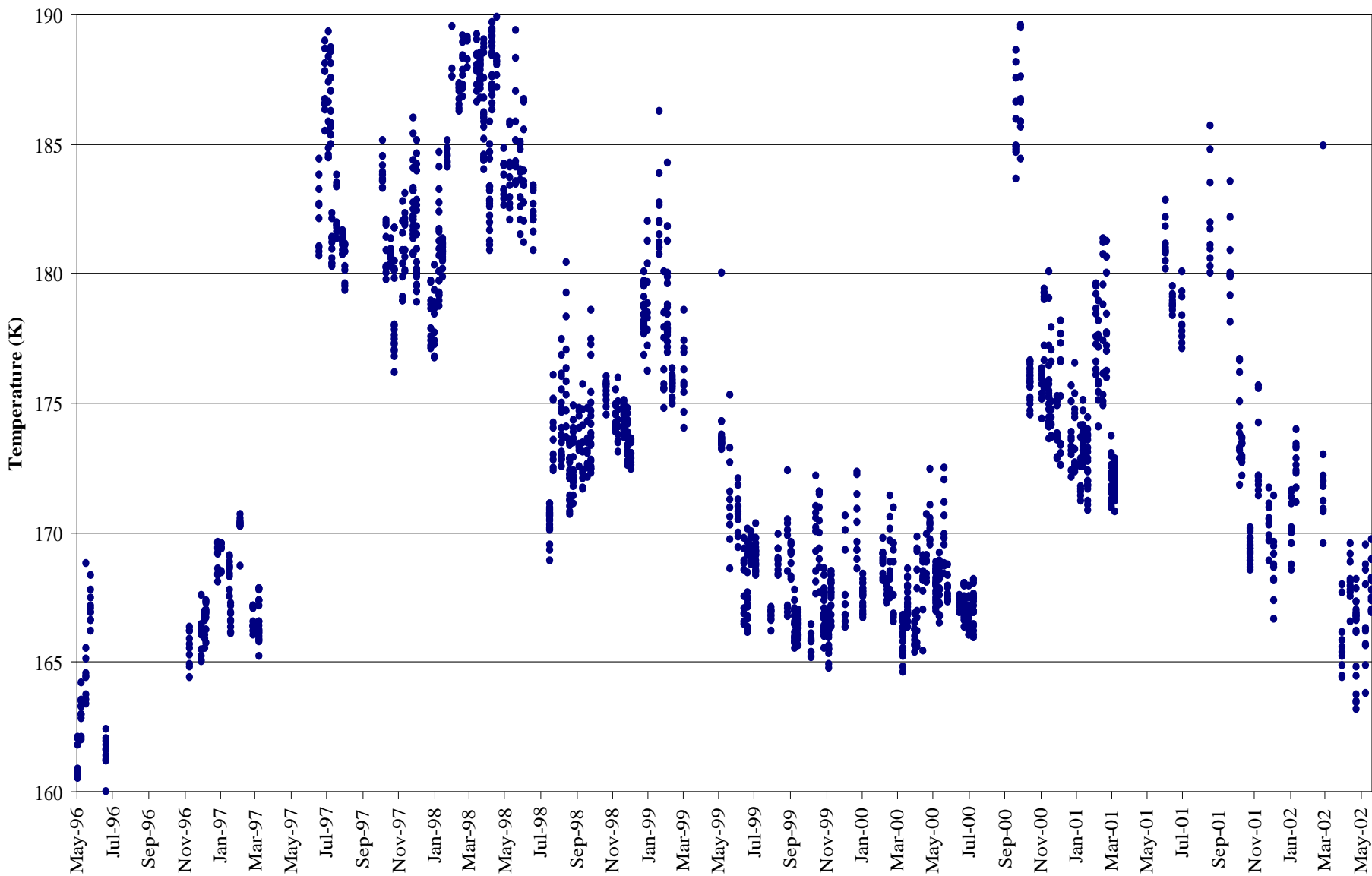
# WVMS3 Noise Diode Calibrations Mauna Loa Hawaii

## Standard Deviation Per Cal

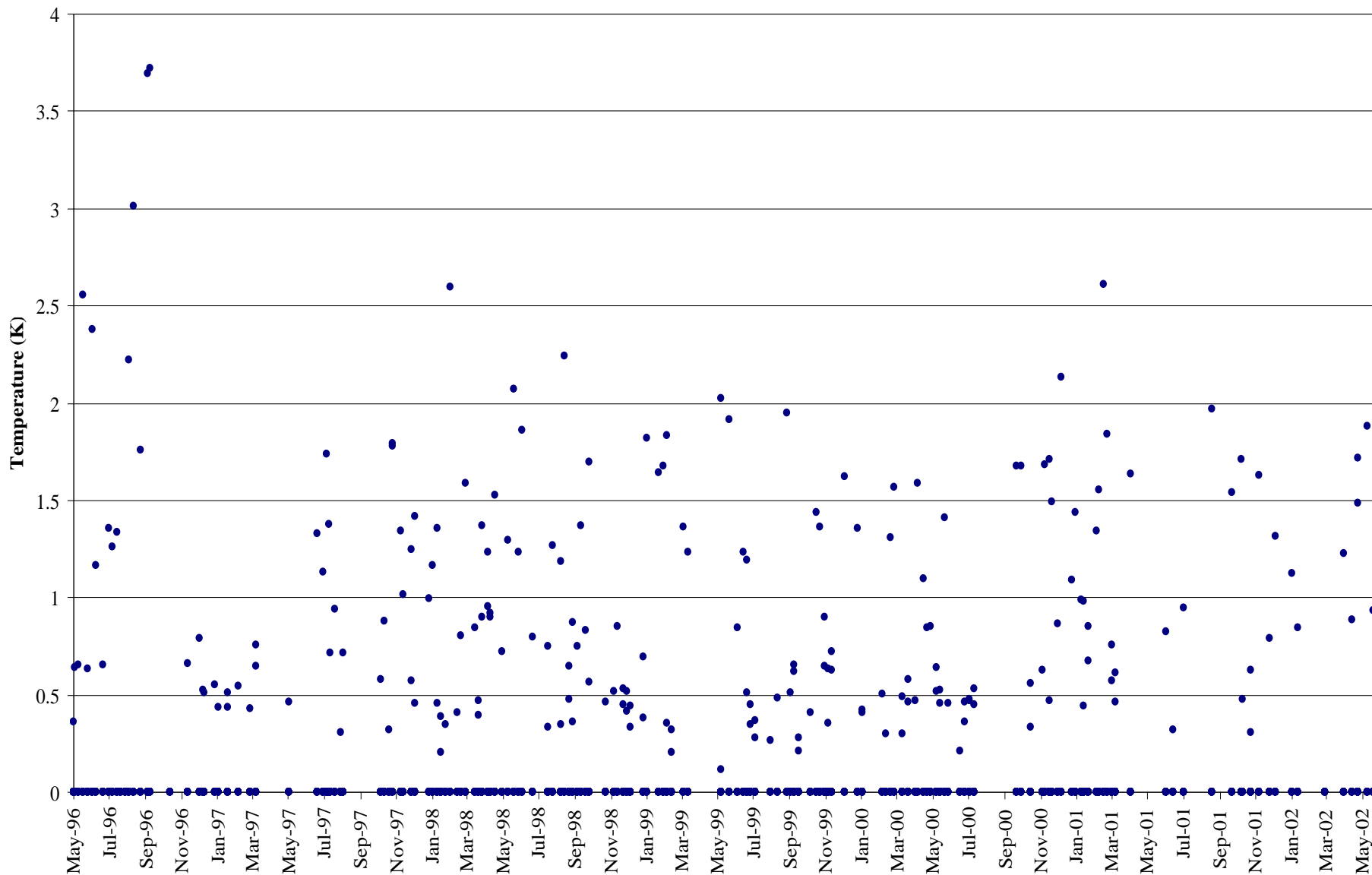




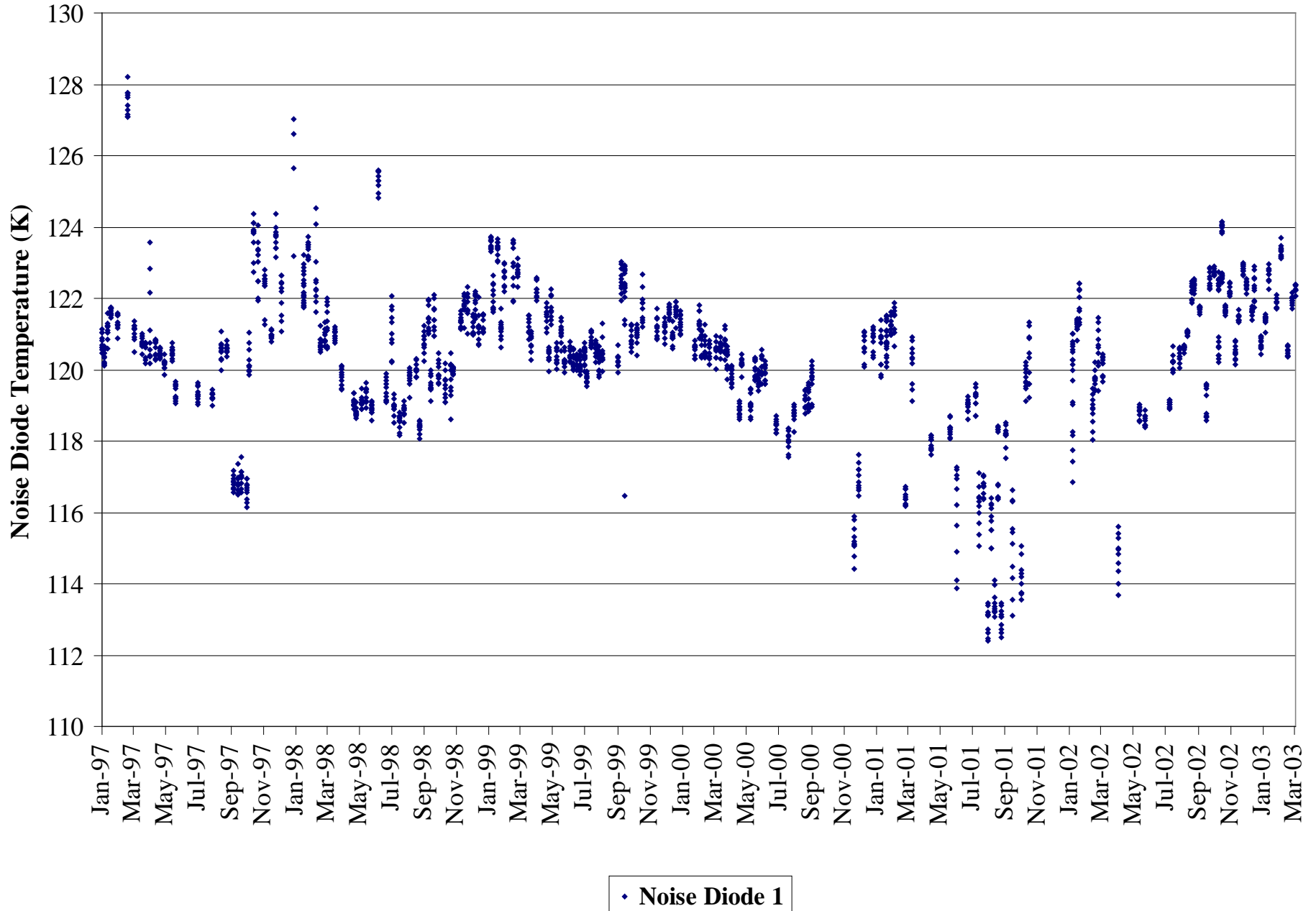
# WVMS3 Receiver Temperature from Calibrations



# WVMS3 Receiver Temperature from Calibrations, $\sigma$ per calibration

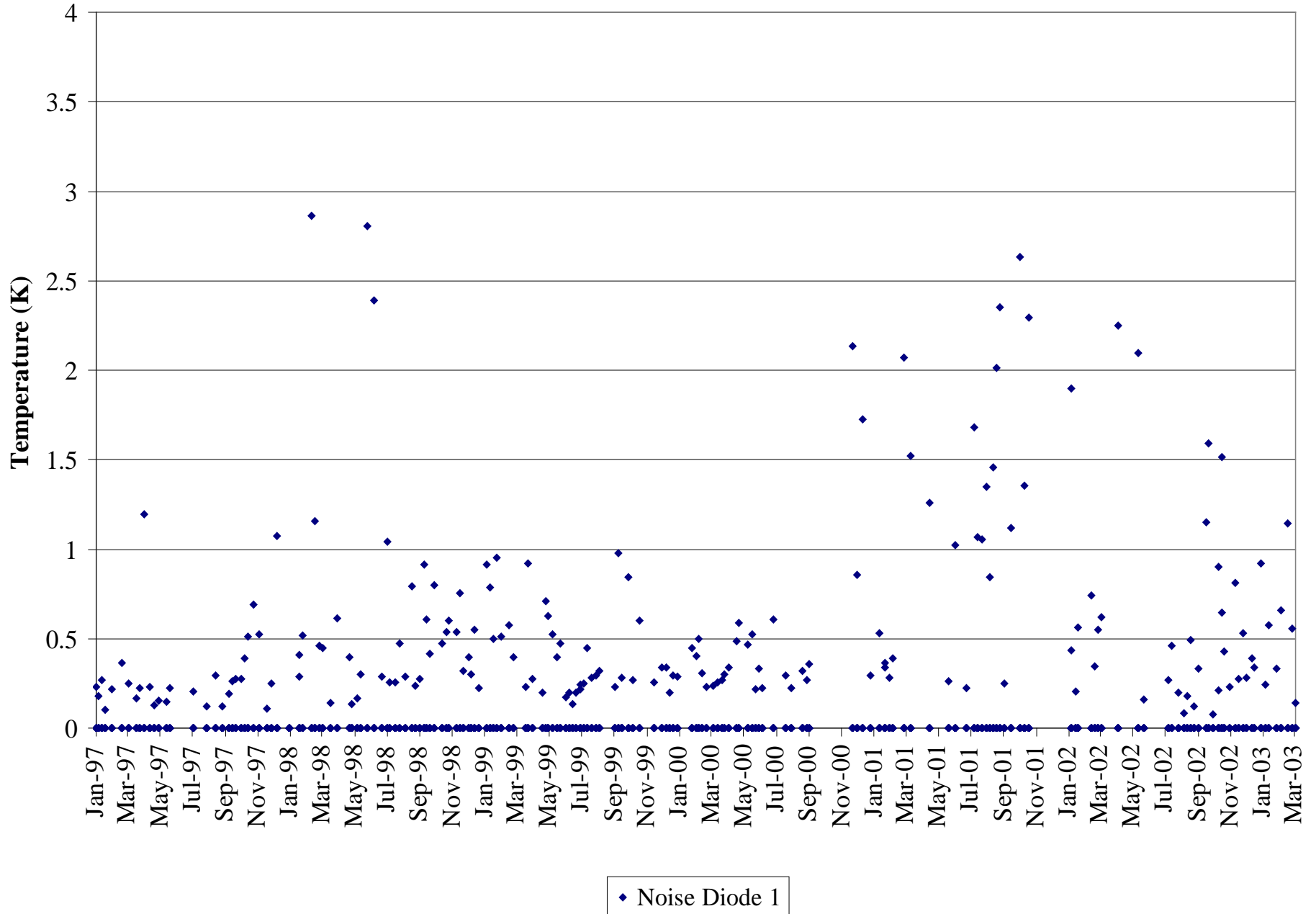


# WVMS1 Noise Diode Calibrations Lauder New Zealand



# WVMS1 Noise Diode Calibrations Lauder New Zealand

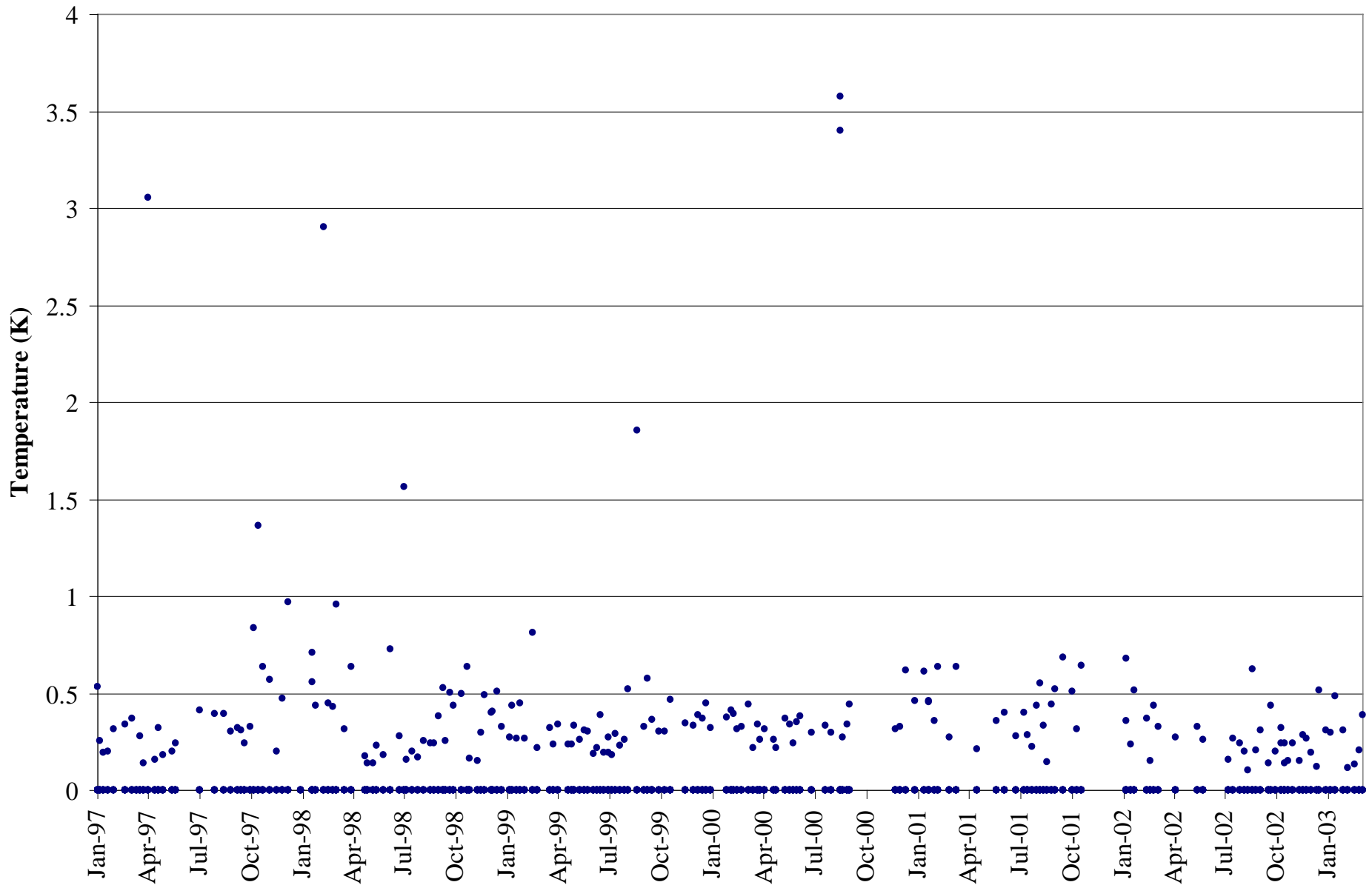
## Standard Deviation Per Cal







# WMS1 Reciever Temperature from Calibrations Lauder New Zealand Standard Deviation Per Cal



# Remote Data Access

- Data is brought back ~weekly from field sites over phone lines using PCAnywhere software
  - No firewall problems
  - Don't have a choice – we're still running DOS programs
- We keep only the tips, the 20 minute averages of the calibrated  $T_{\text{sig}} - T_{\text{ref}}$ , and a few system parameters

# Retrievals

- Optimal estimation based scheme
- 14 MHz channels used only for tips
- Assumed baseline error – Degrades lower altitude retrieval sensitivity
  - Baseline fits were used during a brief period in 1996 at Mauna Loa, when the baselines were unusually bad
  - Spectral resolution is reduced far away from line center, making baseline fits more difficult

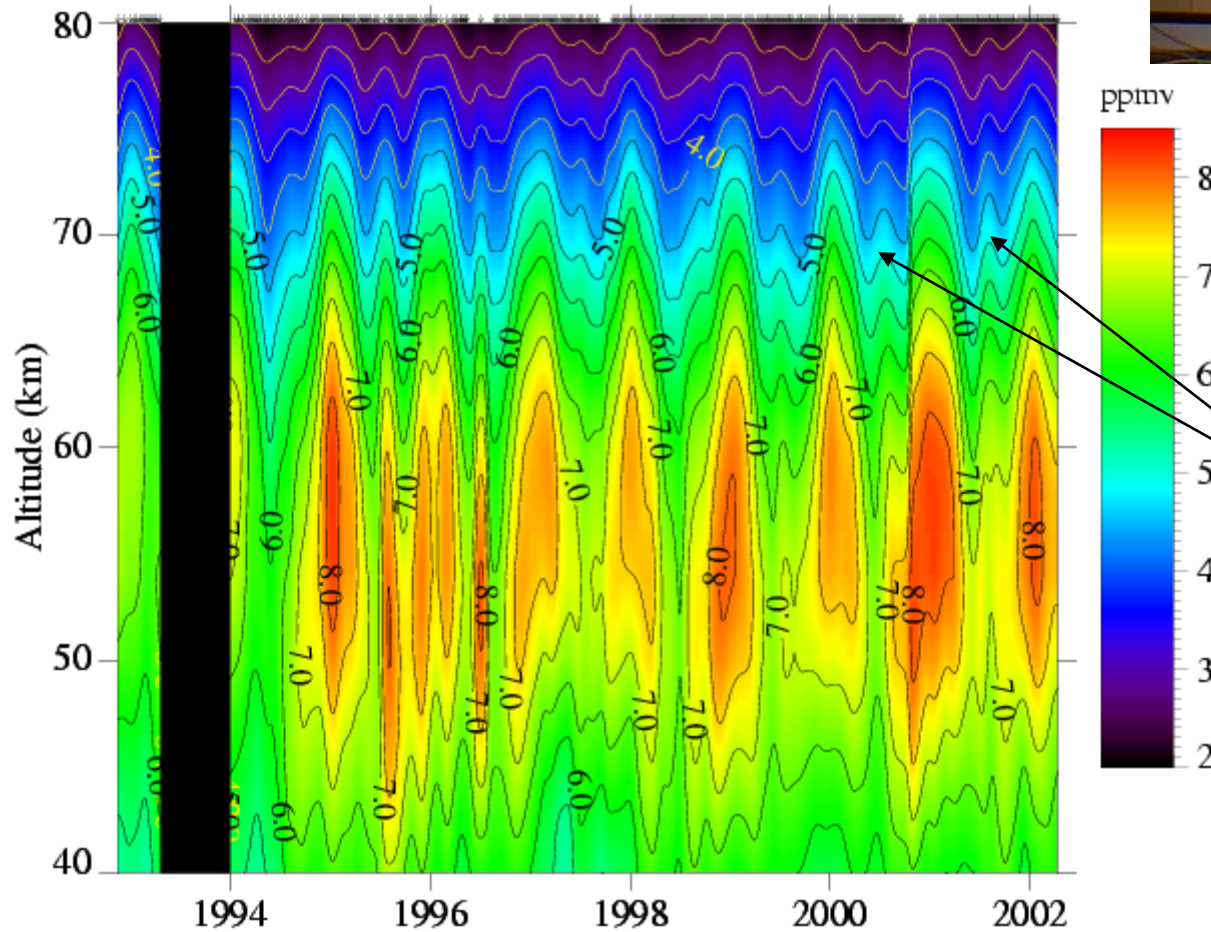
# Retrievals

- Forward model is run all the way to the ground, using tips to provide the tropospheric optical depth (assuming a 2 km scale height for H<sub>2</sub>O)
- Retrievals are generally performed over 500 averaged scans (~1 week)
  - Important to obtain best sensitivity in the mesosphere
  - Reduces a priori influence everywhere – important for trend studies

# Philosophy

- What can a ground-based microwave instrument offer in a world of satellite data?
  - Validation
  - We can fix our instruments!
    - Measurements can be made for as long as spare parts and funding are available –trend studies
  - Continuous observations from one spot
    - Diurnal variations – photochemistry, tides

Data since 1993 from Lauder, New Zealand (45.0S, 169.7E)

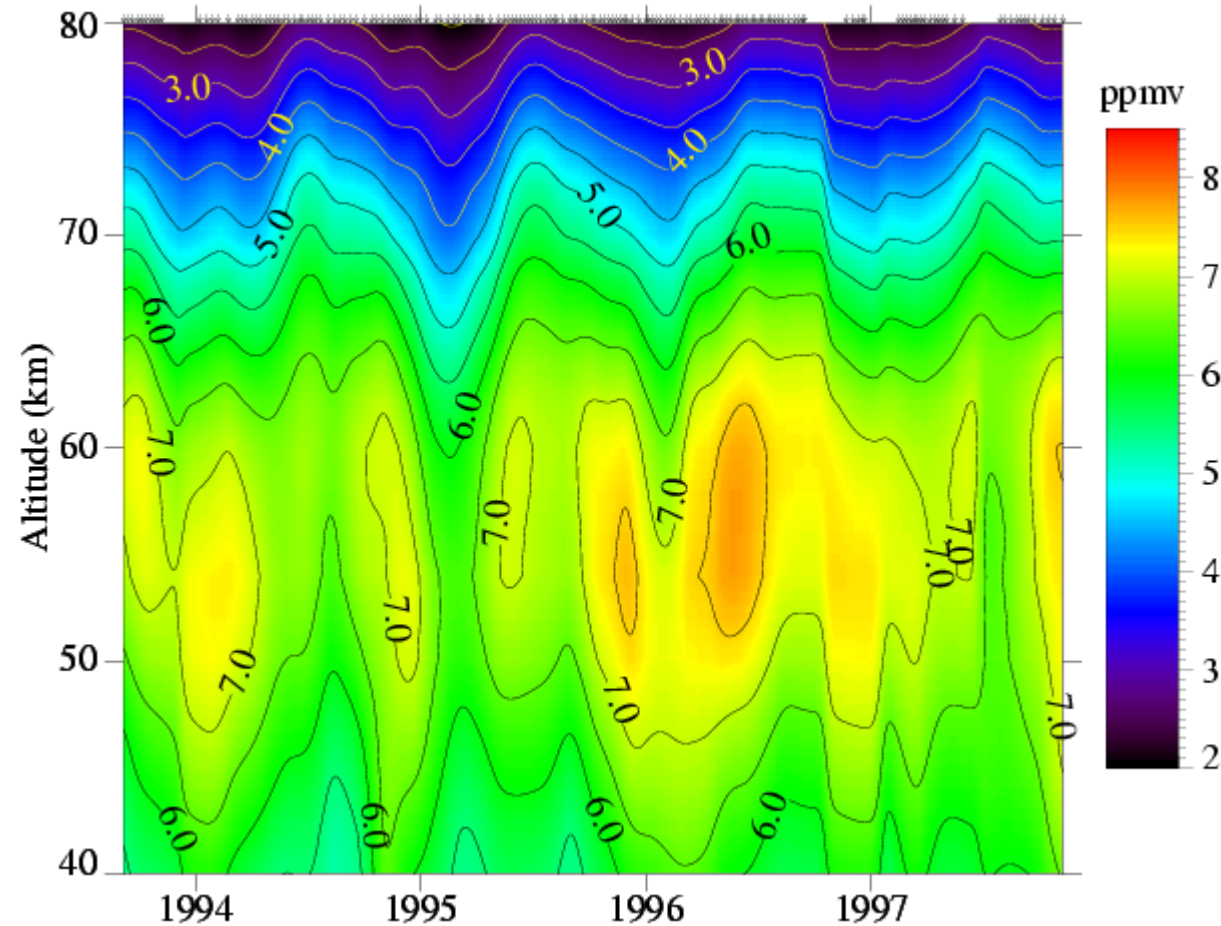


In addition to a large annual cycle, these measurements show a large semi-annual cycle in the upper mesosphere.



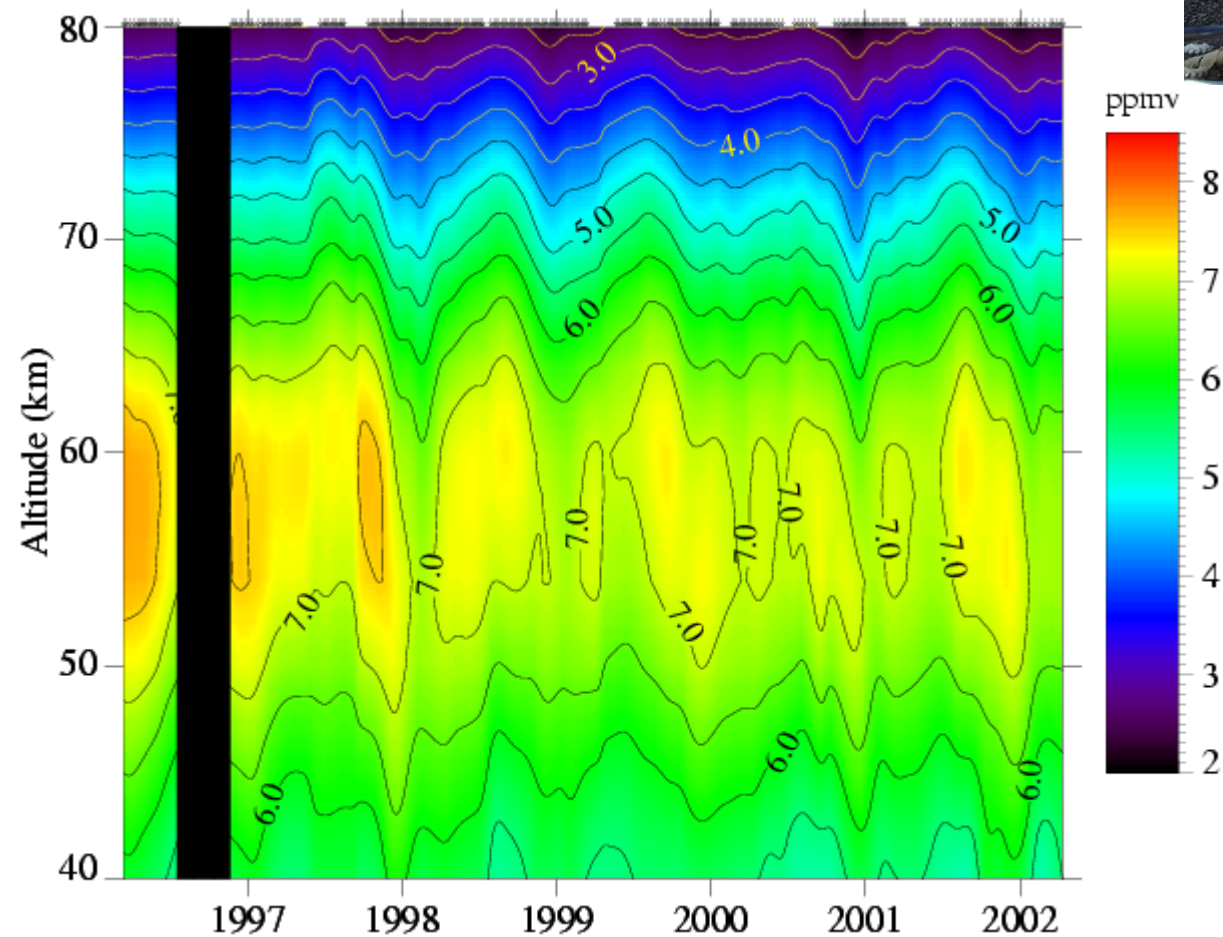
Data from Table Mountain, California  
(34.4N, 242.3E) from 1993-1997

Used to corroborate H<sub>2</sub>O increase  
observed in early 1990s by HALOE





Data since March 1996 from Mauna Loa, Hawaii (19.5N, 204.4E)



Our best dataset  
for future trend  
calculations





# Mechanisms that affect water vapor trends from ~20-60 km

- Tropical tropopause temperature (+0.1 K => ~ +1% water vapor)
- Details of transport from troposphere to stratosphere - Rosenlof
  - Widening of tropical upwelling
  - Changes in ratio of summer/winter net upwelling
- Methane entering the stratosphere
- Middle atmospheric circulation (changes in H<sub>2</sub>O due to changes in CH<sub>4</sub> oxidation – mainly a problem around 40 km)

# Water vapor trend fits

Fit using ( $t$  in years):

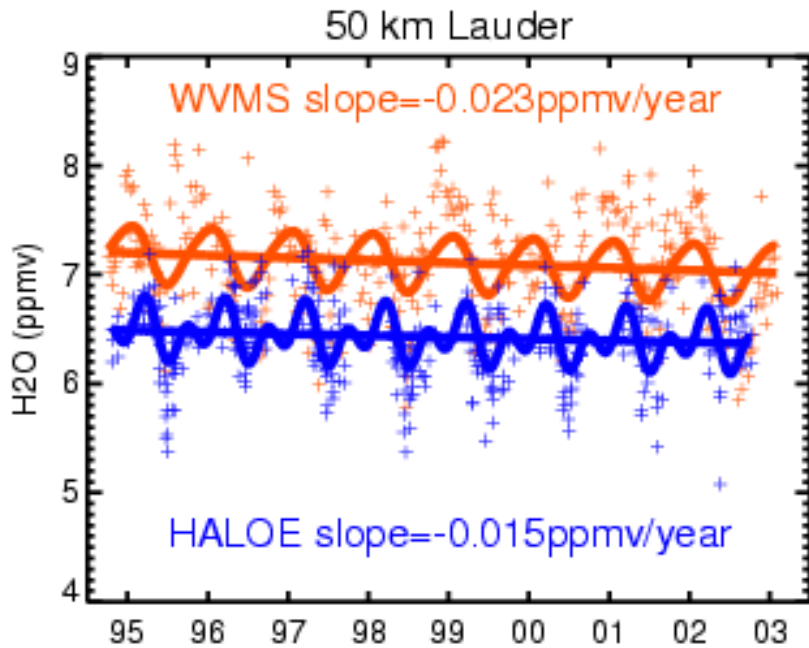
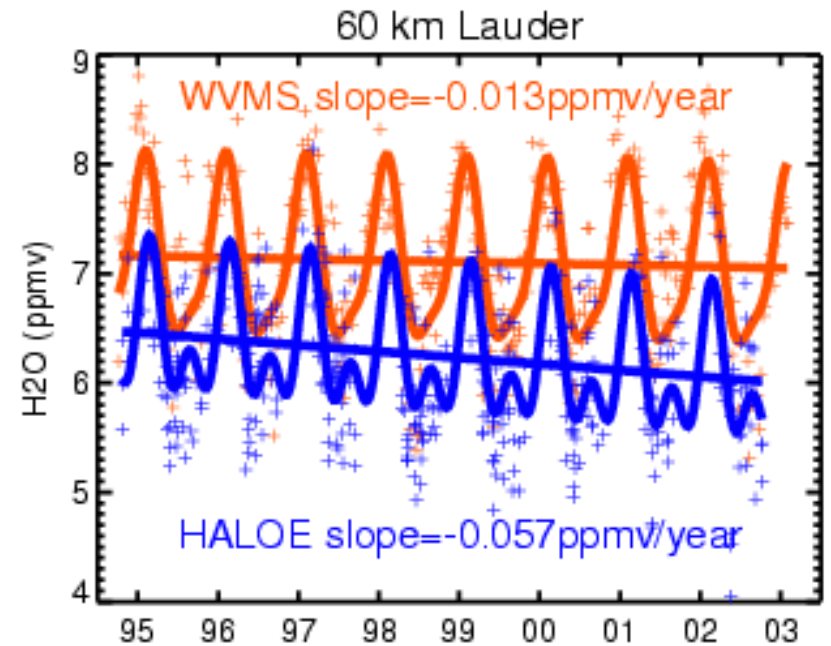
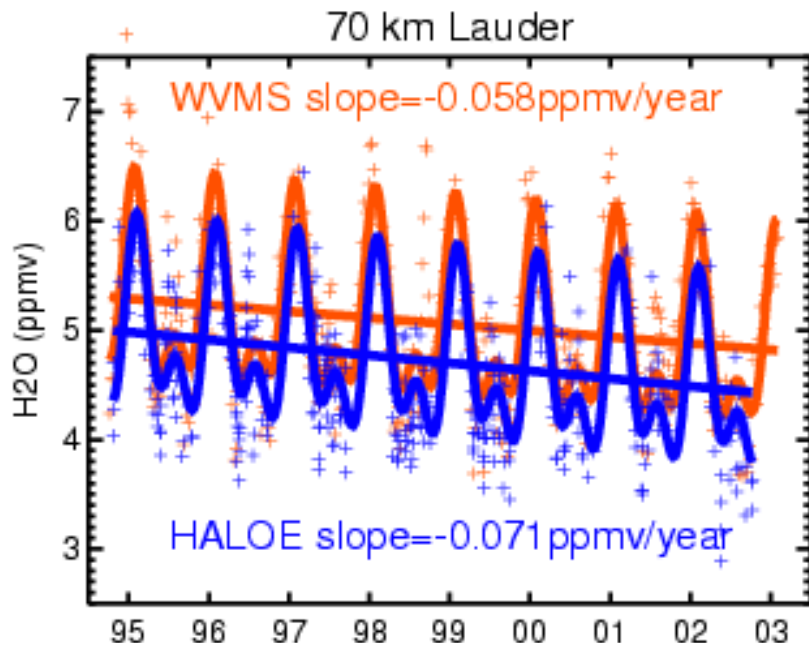
$$f = a_0 +$$

(annual)  $a_1 \sin(2\pi t) + a_2 \cos(2\pi t) +$

(semi-annual)  $a_3 \sin(4\pi t) + a_4 \cos(4\pi t) +$

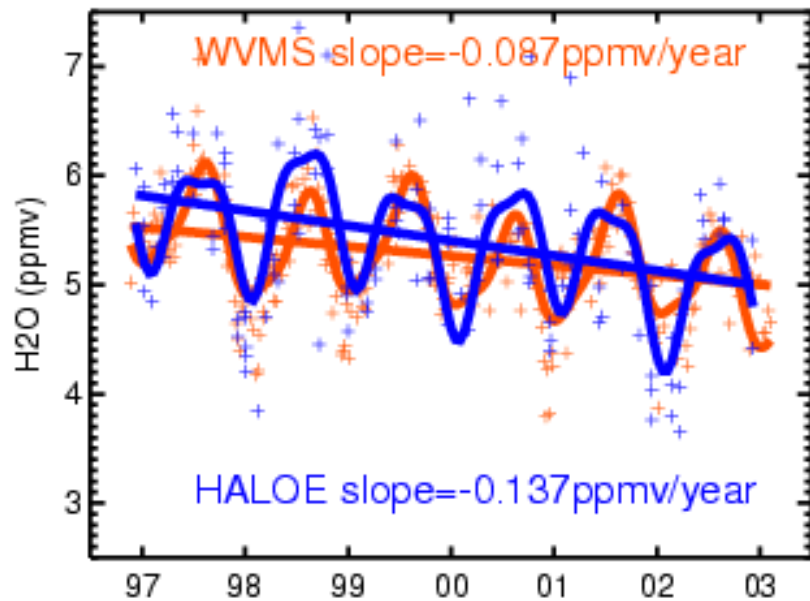
(QBO)  $a_5 \sin((12/27)2\pi t) + a_6 \sin((12/27)2\pi t$

(linear trend)  $+ a_7 t$

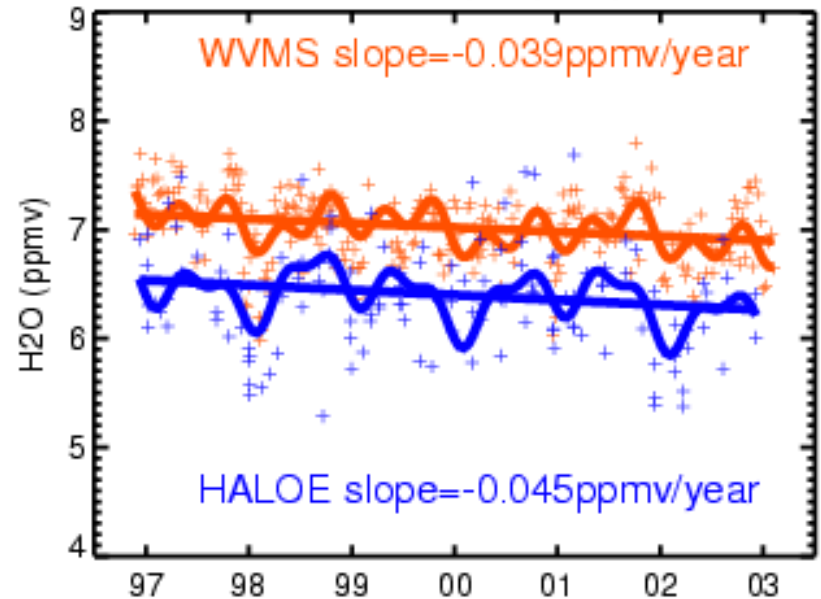


HALOE and WVMS data and best-fits for Lauder

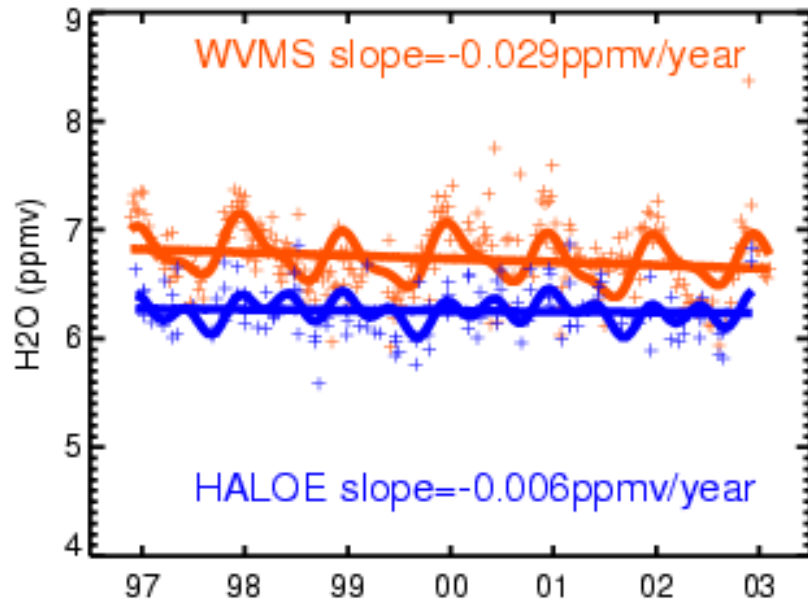
70 km Mauna Loa



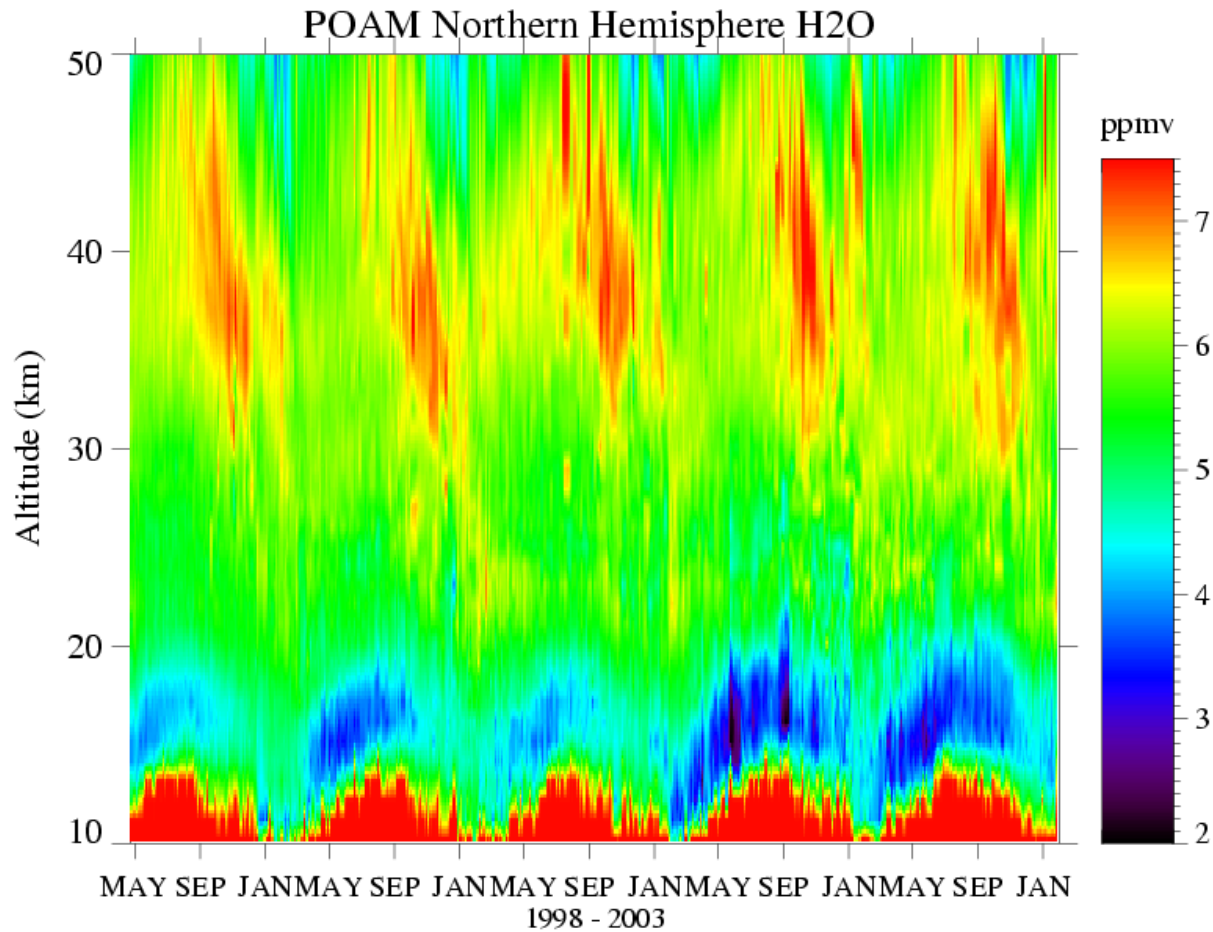
60 km Mauna Loa



50 km Mauna Loa



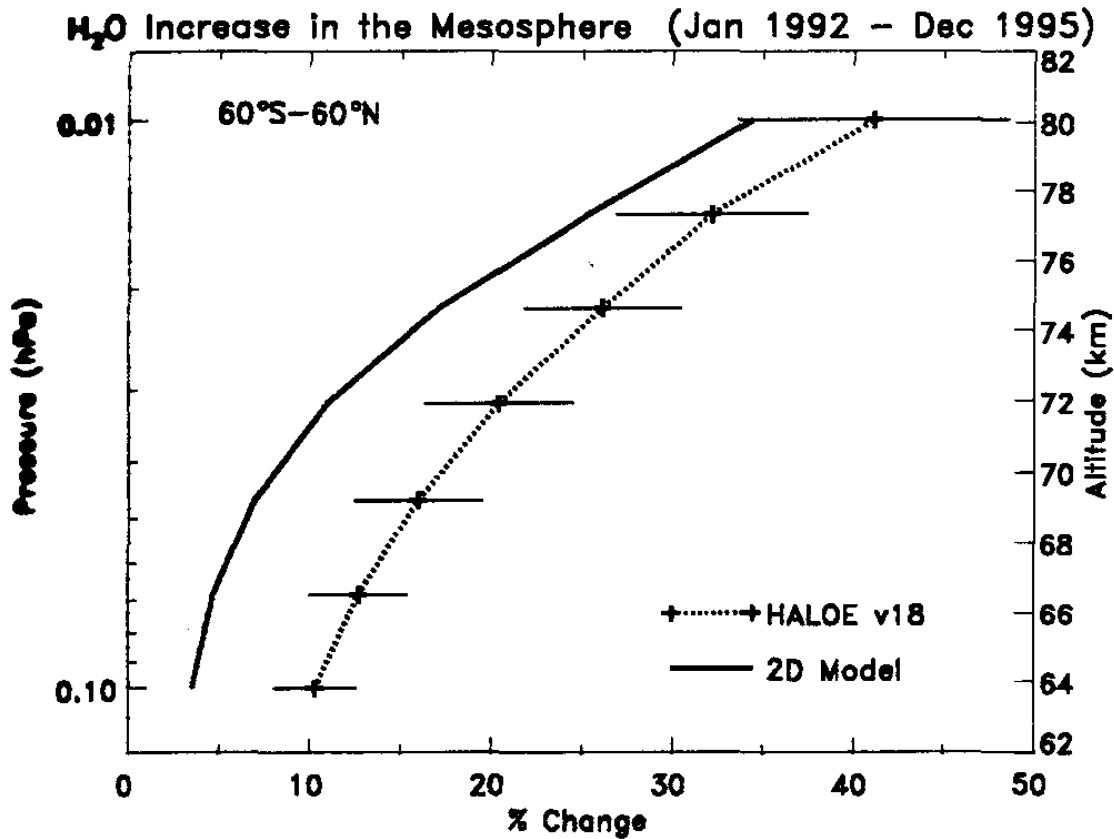
HALOE and WVMS data and best-fits for Mauna Loa



POAM III Northern Hemisphere H<sub>2</sub>O  
measurements since April 1998, covering  
latitudes from 55N-71N

# Summary of linear trend analysis

- WVMS and HALOE both showed a large increase in water vapor in the first half of the 1990s (almost 10%)
- WVMS and coincident HALOE measurements since 1996, and POAM and coincident HALOE measurements since 1998 show smaller linear trends, with roughly an equal number being positive as negative.



**Figure 4.** Height profiles of the 4 year trend estimated from both the model and the HALOE data averaged over 60°S to 60°N.

From Chandra et al., GRL, 1997

# Conclusions

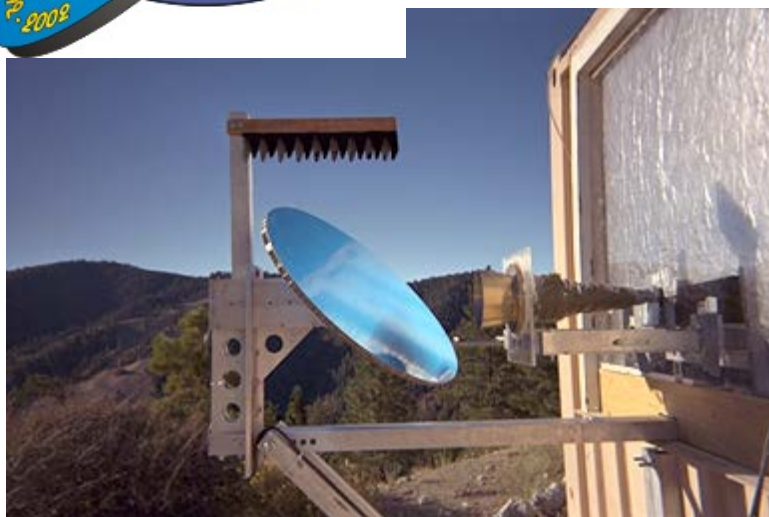
- Large increase in water vapor in the middle atmosphere from 1991-1996 (~10%).
- No clear trend from 1996-present.
- Some indication of an increase over last 2-3 years.
- Complicated pattern of water vapor variations gives an increase of ~0.5 ppmv (~1%/year) for the HALOE mission (1992-present).
- A 1%/year increase is very similar to the annual increase quoted by Rosenlof et al. [2001] for the years 1954-2001, and by Oltmans et al. [2000] for the years 1964-2000.





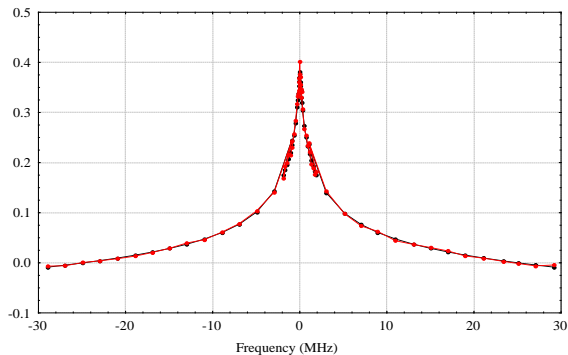
# WVMS

## Vapor Millimeter-wave

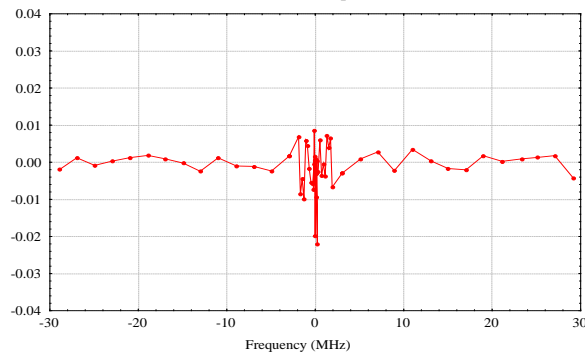


- Part of the international Network for the Detection of Stratospheric Change (NDSC)
- Continuous measurements of water vapor in the stratosphere and mesosphere to determine seasonal variability and interannual changes
- Ground based microwave spectroscopy to measure the pressure broadened 22.2 GHz water vapor line
- Retrieve mixing ratios using optimal estimation

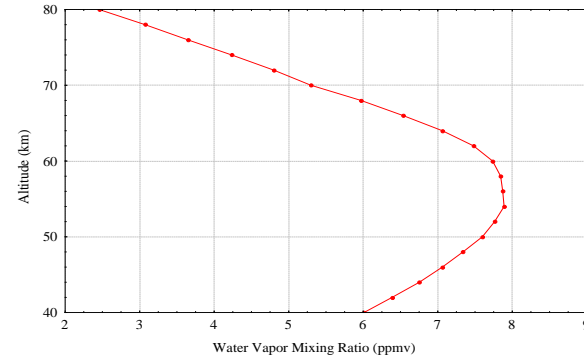
WVMS3 Mauna Loa April 7, 1996



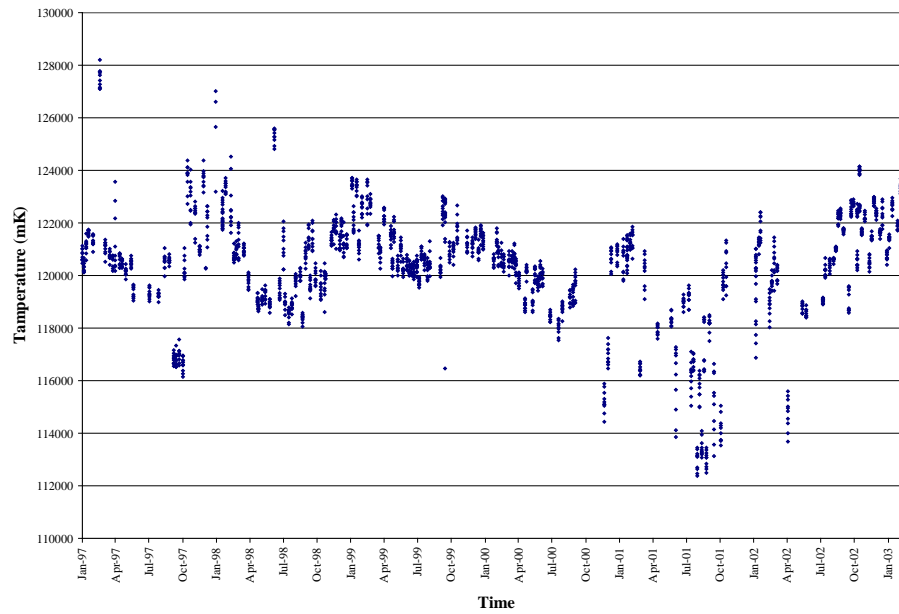
WVMS3 Mauna Loa April 7, 1996



WVMS3 Mauna Loa April 7, 1996



WVMS1 Calibration Data



WVMS1 Receiver Temperature from Calibrations

